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2 **Assessing climate model projections: state of the art and philosophical reflections**

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Abstract

The present paper draws on climate science and the philosophy of science in order to evaluate climate-model-based approaches to assessing climate projections. We analyze the difficulties that arise in such assessment and outline criteria of adequacy for approaches to it. In addition, we offer a critical overview of the approaches used in the IPCC working group one fourth report, including the confidence building, Bayesian and likelihood approaches. Finally, we consider approaches that do not feature in the IPCC reports, including three approaches drawn from the philosophy of science. We find that all available approaches face substantial challenges, with IPCC approaches having as a primary source of difficulty their goal of providing probabilistic assessments.

83

84 **1. Introduction**

85 The climate system is the system of processes that underlie the behavior of
 86 atmospheric, oceanic and cryospheric phenomena such as atmospheric temperature,
 87 precipitation, sea-ice extent and ocean salinity. Climate models are designed to
 88 simulate the seasonal and longer term behavior of the climate system. They are
 89 mathematical, computer implemented representations that comprise two kinds of
 90 elements. They comprise basic physical theory – e.g., conservation principles such as
 91 conservation of momentum and heat – that is used explicitly to describe the evolution
 92 of some physical quantities – e.g., temperature, wind velocity and properties of water
 93 vapor. Climate models also comprise parameterizations. Parameterizations are
 94 substitutes for explicit representations of physical processes, substitutes that are used
 95 where lack of knowledge and/or limitations in computational resources make explicit
 96 representation impossible. Individual cloud formation, for example, typically occurs
 97 on a scale that is much smaller than global climate model (GCM) resolution and thus
 98 cannot be explicitly resolved. Instead, parameterizations capturing assumed
 99 relationships between model grid-average quantities and cloud properties are used.

100 The basic theory of a climate model can be formulated using equations for the
 101 time derivatives of the model's state vector variables, x_i , $i = 1, \dots, n$, as is
 102 schematically represented by

$$103 \quad \frac{\partial x_i}{\partial t} = F_i(x_1, \dots, x_n, y_1, \dots, y_n, t) + G_i(t) \quad (1)$$

104 In Eq. (1), t denotes time, the functions G_i represent external forcing factors
 105 and how these function together to change the state vector quantities, and the F_i
 106 represent the many physical, chemical and biological factors in the climate system and
 107 how these function together to change the state vector quantities. External forcing

108 factors – e.g., greenhouse gas concentrations, solar irradiance strength, anthropogenic
109 aerosol concentrations and volcanic aerosol optical depth – are factors that might
110 affect the climate system but that are, or are treated as being, external to this system.

111 The x_i represent those quantities the evolution of which is explicitly described
112 by basic theory, that is the evolution of which is captured by partial time derivatives.
113 The y_i represent quantities that are not explicitly described by basic theory. So these
114 variables must be treated as functions of the x_i , i.e., the y_i must be parameterized. In
115 this case, the parameterizations are schematically represented in Eq. (2).

$$116 \quad y_i = H_i(x_1, \dots, x_n) \quad (2)$$

117 Given initial conditions $x_i(t_0)$ at time $t = t_0$ and boundary conditions, the climate
118 model calculates values of the state vector at a later time $t = t_f$ in accordance with
119 Eq. (1).

120 Climate models play an essential role in identifying the causes of climate
121 change and in generating projections. Projections are conditional predictions of
122 climatic quantities. Each projection tells us how one or more such quantities would
123 evolve were external forcing to be at certain levels in the future. Some approaches to
124 assessing projections derive projections, and assess their quality, at least partly
125 independently of climate models. They might, for example, use observations to decide
126 how to extend simulations of present climate into the future (Stott et al., 2006) or
127 derive projections from, and assess them on the basis of, observations (Bentley, 2010;
128 Siddall et al., 2010). We focus on climate-model-based assessment. Such assessment
129 is of the projections of one or more climate models and is assessment in which how
130 good models are in some respect or another is used to determine projection quality. A
131 climate model projection (CMP) quality is a qualitative or quantitative measure, such
132 as a probability, that is indicative of what we should suppose about CMP accuracy.

133 It is well recognized within the climate science community that climate-
134 model-based assessment of projection quality needs to take into account the effects of
135 climate model limitations on projection accuracy (Randall et al., 2007; Smith, 2006;
136 Stainforth et al., 2007a). Following Smith (2006) and Stainforth (2007a), we
137 distinguish between the following main types of climate model limitations:

138 (a) External forcing inaccuracy – inaccuracy in a model's representation of
139 external forcing, that is in the G_i in Eq. (1).

140

141 (b) Initial condition inaccuracy – inaccuracy in the data used to initialize
142 climate model simulations, that is in the $x_i(t_0)$.

143

144 (c) Model imperfection – limitations in a model's representation of the climate
145 system or in our knowledge of how to construct this representation,
146 including:

147

148 1. Model parameterization limitations – limitations in our knowledge of
149 what the optimal or the appropriate parameter values and parameterization
150 schemes for a model are. This amounts, in the special case where
151 parameterizations are captured by Eq. (2), to limitations in our knowledge
152 of which functions H_i one should include from among available
153 alternatives.

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155 2. Structural inadequacy – inaccuracy in how a model represents the
156 climate system which cannot be compensated for by resetting model
157 parameters or replacing model parameterizations with other available
158 parameterization schemes. Structural inaccuracy in Eq. (1) is manifested
159 in an insufficient number of variables x_i and y_i as well as in the need for
160 new functions of these variables.

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162 Parameterization limitations are illustrated by the enduring uncertainty about climate
163 sensitivity and associated model parameters and parameterization schemes. A
164 relatively recent review of climate sensitivity estimates underscores the limited ability
165 to determine its upper bound as well as the persistent difficulty in narrowing its likely
166 range beyond 2 to 4.5 °C (Knutti and Hegerl, 2008). The 21 GCMs used by Working
167 Group One of the IPCC fourth report (WG1 AR4) illustrate structural inadequacy.
168 These sophisticated models are the models of the World Climate Research
169 Programme's Coupled Model Intercomparison Project phase 3 (CMIP3) (Meehl et al.,

170 2007a). Some important sub-grid and larger than grid phenomena that are relevant to
171 the evolution of the climate system are not accurately represented by these models,
172 some are only represented by a few of the models and some are not represented at all.
173 Parameterization of cloud formation, for example, is such that even the best available
174 parameterizations suffer from substantial limitations (Randall et al., 2003). None of
175 the models represent the carbon cycle, only some represent the indirect aerosol effect
176 and only two represent stratospheric chemistry (CMIP3, 2007). The models also omit
177 many of the important effects of land use change (Mahmood et al., 2010; Pielke,
178 2005). Many of their limitations, e.g., the limited ability to represent surface heat
179 fluxes as well as sea ice distribution and seasonal changes, are the result of a
180 combination of structural inadequacy and parameterization limitations (Randall et al.,
181 2007, p. 616). CMIP3 simulations illustrate initial condition inaccuracy. Due to
182 constraints of computational power and to limited observations, these simulations start
183 from selected points of control integrations rather than from actual observations of
184 historical climate (Hurrell et al., 2009).

185 The most ambitious assessments of projection quality, and these are primarily
186 climate-model-based assessments, are those of WG1. The first three WG1 reports rely
187 primarily on the climate-model-based approach that we will call the confidence
188 building approach. This is an informal approach that aims to establish confidence in
189 models, and thereby in their projections, by appealing to models' physical basis and
190 success at representing observed and past climate. In the first two reports, however,
191 no uniform view about what confidence in models teaches about CMP quality is
192 adopted (IPCC 1990; IPCC 1996). The summary for policymakers in the WG1
193 contribution to the IPCC first assessment report, for example, qualifies projections
194 using diverse phrases such as 'we predict that', 'confidence is low that' and 'it is likely

195 that' (IPCC 1990). A more systematic view is found in WG1's contribution to the
196 third IPCC assessment report (WG1 TAR). It made use of a guidance note to authors
197 which recommends that main results be qualified by degrees of confidence that are
198 calibrated to probability ranges (Moss and Schneider, 2000). The summary for
199 policymakers provided by WG1 TAR does assign projections such degrees of
200 confidence. It expresses degrees of confidence as degrees of likelihood and takes, e.g.,
201 'very likely' to mean having a chance between 90 and 99 %, and 'likely' to mean
202 having a chance between 66 % and 90 %. The chapter on projections of future climate
203 change, however, defines degrees of confidence in terms of agreement between
204 models. A very likely projection, for example, is defined (roughly) as one that is
205 physically plausible and is agreed upon by all models used (IPCC 2001).

206 WG1 AR4's assessment of projection quality has two stages. First, confidence
207 in models is established as in previous reports. This is mostly achieved in Chapter 8 –
208 which describes, among other things, successful simulations of natural variability
209 (Randall et al., 2007) – and in chapter 9 – which focuses on identifying the causes of
210 climate change, but also characterizes model successes at simulating 20th century
211 climate change (Hegerl et al., 2007). The second stage is carried out in Chapter 10 –
212 which provides WG1 AR4's global projections (Meehl et al., 2007b) – and Chapter 11
213 – which focuses on regional projections (Christensen et al., 2007). In these chapters,
214 expert judgment is used to assign qualities to projections given established confidence
215 in models and the results of formal, probabilistic projection assessment (Meehl et al.,
216 2007b). WG1 AR4 is the first WG1 report that makes extensive use of formal
217 assessment, though it recognizes that such approaches are in their infancy
218 (Christensen et al., 2007; Randall et al., 2007). Both climate-model-based and partly
219 climate-model-independent formal approaches are used.

220 Although WG1 AR4 assesses models using degrees of confidence, it does not
221 assess projections in these terms. Nor does it equate projection likelihoods with
222 degrees of agreement among models. It does, however, implement the advice to
223 provide probabilistically calibrated likelihoods of projections (IPCC 2005). For
224 example, unlike WG1 TAR, WG1 AR4 provides explicit likelihood estimates for
225 projected ranges of global mean surface temperature (GMST) changes. It estimates
226 that the increase in GMST by the end of the century is likely to fall within -40 to +60
227 % of the average GCM warming simulated for each emission scenario and provides
228 broader uncertainty margins than the GCM ensemble in particular because GCMs do
229 not capture uncertainty in the carbon cycle (Fig. 2).

230 The sophistication of WG1 AR4's assessments was enabled by the increasing
231 ability to use multi-GCM and perturbed physics GCM ensembles. Thus, while WG1's
232 first two reports relied on simple models to produce long term GMST projections,
233 WG1 TAR and WG1 AR4 relied primarily on state-of-the-art GCM ensembles to
234 assess these and other projections. WG1 AR4 nevertheless still relied on simpler
235 models, including intermediate complexity and energy balance models (Randall et al.,
236 2007).

237 In this review, we provide a critical discussion of the (climate-model-based)
238 approaches to assessing projection quality relied on in WG1 AR4 and more recent
239 work by climate scientists. In doing so, we build on the substantial climate science
240 literature, including WG1 AR4 itself. We, however, extend this literature using the
241 perspective of the philosophy of science. Our discussion does focus more than climate
242 scientists themselves tend to on precisely why assessing projection quality is difficult,
243 on what is required of an adequate approach to such assessment and on the limitations
244 of existing approaches. We, nevertheless, also address some of the practical concerns

245 of climate scientists. We outline three views of how to assess scientific claims that are
246 drawn from the philosophy of science and consider how they might further assist in
247 assessing projection quality. Important issues that space does not allow us to address
248 are the special difficulties that assessment of regional projection quality raises. An
249 issue that deserves more attention than we have given it is that of how uncertainty
250 about data complicates assessing projection quality.

251 We begin (Section 2) by considering what kinds of qualities should be
252 assigned to projections, especially whether probabilistic qualities should be assigned.
253 We then (Section 3) discuss why assessing projection quality is difficult and outline
254 criteria for adequate approaches to doing so. Using these criteria, we proceed to
255 discuss (Sections 4–7) the approaches that were used in WG1 AR4, namely the
256 confidence building, the subjective Bayesian and the likelihood approaches. Finally
257 (Section 8), we discuss approaches that are not used, or are not prominent in, WG1
258 AR4, including the possibilist and three philosophy-of-science-based approaches.

259

260 **2. Probabilistic and non-probabilistic assessment**

261 Probabilistic assessment of projection quality will here be taken to include assigning
262 probabilities or informative probability ranges to projections or projection ranges.
263 Such assessment has been argued for on the ground that it is better suited to handling
264 the inevitable uncertainty about projections than deterministic assessments are
265 (Raisanen and Palmer, 2001). But philosophers of science, computer scientists and
266 others point out that probabilities fail to represent uncertainty when ignorance is deep
267 enough (Halpern, 2003; Norton, 2011). Assigning a probability to a prediction
268 involves, given standard probability frameworks, specifying the space of possible
269 outcomes as well as the chances that the predicted outcomes will obtain. These,

270 however, are things we may well be uncertain about given sufficient ignorance. For
271 example, we might be trying to assess the probability that a die will land on '6' when
272 our information about the kind and bias of the die is limited. We might have the
273 information that it can exhibit the numerals '1', '6' and '8' as well as the symbol '*', but
274 not have any information about what other symbols might be exhibited or, beyond the
275 information that '6' has a greater chance of occurring than the other known symbols,
276 the chances of symbols being exhibited. The die need not be a six sided die. In such
277 circumstances, it appears that assigning a probability to the outcome '6' will
278 misrepresent our uncertainty.

279 Assigning probability ranges and probabilities to ranges can face the same
280 difficulties as assigning probabilities to single predictions. In the above example,
281 uncertainty about the space of possibilities is such that it would be inappropriate to
282 assign the outcome '6' a range that is more informative than the unhelpful 'somewhere
283 between 0 and 1'. The same is true about assigning the range of outcomes '1', '6' and
284 '8' a probability.

285 One might suggest that, at least when the possible states of a system are
286 known, we should apply the principle of indifference. According to this principle,
287 where knowledge does not suffice to decide between possibilities in an outcome
288 space, they should be assigned equal probabilities. Some work in climate science
289 acknowledges that this principle is problematic, but suggests that it can be applied
290 with suitable caution (Frame et al., 2005). Most philosophers argue that the principle
291 should be rejected (Strevens, 2006a). We cannot know that the principle of
292 indifference will yield reliable predictions when properly applied (North, 2010). If,
293 for example, we aim to represent complete ignorance of what value climate sensitivity
294 has within the range 2 to 4.5 °C, it is natural to assign equal probabilities to values in

295 this range. Yet whether doing so is reliable across scenarios in which greenhouse
296 gasses double depends on what climate sensitivity actually tends to be across such
297 scenarios and it is knowledge of this tendency that is, given the assumed ignorance,
298 lacking. Further, we can only define a probability distribution given a description of
299 an outcome space and there is no non-arbitrary way of describing such a space under
300 ignorance (Norton, 2008; Strevens, 2006a). What probability should we assign to
301 climate sensitivity's being between 2 and 4 °C, given complete ignorance within the
302 range 2 to 6 °C? 50 % is the answer, when the outcome space is taken to be the given
303 climate sensitivity range and outcomes are treated as equiprobable. But other answers
304 are correct if alternative outcome spaces are selected, say if the outcome space is
305 taken to be a function not just of climate sensitivity but also of feedbacks upon which
306 climate sensitivity depends. And in the supposed state of ignorance about climate
307 sensitivity, we will not have a principled way of selecting a single outcome space.

308 Although the case of the die is artificial, our knowledge in it does share some
309 features with our knowledge of the climate system. We are, for example, uncertain
310 about what possible states the climate system might exhibit, as already stated in the
311 case of climate sensitivity. A central question in what follows is to what extent our
312 ignorance of the climate system is such that probabilistic assessment of projection
313 quality is inappropriate.

314 Acknowledging that probabilistic assessment is inappropriate in some case is
315 by no means then to give up on assessment. Assigning non-probabilistic qualities can
316 commit us to less than assigning probabilities or probability ranges and thus can better
317 represent uncertainty. Judging that it is a real possibility that climate sensitivity is 2
318 °C does not require taking a position on the full range of climate sensitivity. Nor need
319 rankings of climate sensitivities according to plausibility do so. Other non-

320 probabilistic qualities the assignment of which is less demanding than that of
321 probabilities or probability ranges are sets of probability ranges and the degree to
322 which claims have withstood severe tests (see Halpern (2003) for a discussion, and
323 formal treatment, of a variety of non-probabilistic qualities. We discuss severe-test-
324 based and real-possibility-based assessments in sections 8.4 and 8.1 respectively).

325

326 **3. Why is assessing projection quality difficult?**

327 Projections, recall, are predictions that are conditional on assumptions about external
328 forcing. So errors in assumptions about external forcing are not relevant to assessing
329 projection quality. Such assessment need only take into account the effects of initial
330 condition inaccuracy and model imperfection. In the present section, we consider why
331 these kinds of limitations make assessing projection quality difficult. This question is
332 not answered just by noting that climate models have limitations. Scientific models
333 are in general limited, but it is not generally true that assessing their predictions is a
334 serious problem. Consider standard Newtonian models of the Earth-Sun system. Such
335 models suffer from structural inadequacy. They represent the Earth and the Sun as
336 point masses. Moreover, they tell us that the Earth and the Sun exert gravitational
337 forces on each other, something that general relativity assures us is not strictly true.
338 Still, assessing to what extent we can trust the predictions these models are used to
339 generate is something we typically know how to do.

340

341 **3.1 Initial condition inaccuracy and its impact on assessing projections**

342 We begin by considering the difficulties associated with initial condition error. Work
343 in climate science emphasizes the highly nonlinear nature of the climate system (Le
344 Treut et al., 2007; Rial et al., 2004), a nature that is reflected in the typically nonlinear

345 form of the F_i in Eq. (1). Nonlinear systems are systems in which slight changes to
346 initial conditions can give rise to non-proportional changes of quantities over time
347 (Lorenz, 1963). This high sensitivity can make accurate prediction inherently difficult.
348 Any errors in simulations of highly nonlinear systems, including even minor errors in
349 initial condition settings, might be multiplied over time quickly. The high sensitivity
350 to initial conditions also, as climate scientists note, threatens to make assessing
351 prediction quality difficult. The way in which error grows over time in such systems
352 cannot be assumed to be linear and might depend on how the system itself develops
353 (Palmer, 2000; Palmer et al., 2005).

354 However, how serious a problem sensitivity to initial conditions is for
355 assessing projection quality is not a straightforward matter. The known inaccuracy in
356 model initial condition settings means that high sensitivity of the evolution of climatic
357 quantities to initial conditions might be important. Yet, a climatic quantity the
358 evolution of which is going to be highly nonlinear at one temporal scale may continue
359 to exhibit approximately linear evolution on another such scale. Greenland ice volume
360 may, for example, evolve linearly in time over the coming few decades but
361 nonlinearly over more than three centuries (Lenton et al., 2008). If this is so,
362 nonlinearity will only be a limited obstacle to assessing projections of Greenland ice
363 volume. More generally, whether, and to what extent, a climatic process is nonlinear
364 will depend on the desired projection accuracy, the quantity of interest, the actual
365 period and region of interest and the temporal and spatial scale of interest (IPCC
366 2001). Thus, whether the highly nonlinear behavior of the climate system is a problem
367 for assessing projection quality will have to be determined on a case by case basis.

368

369 **3.2 Tuning and its impact on assessing projections**

370 Further features of climate modeling complicate determining the impact of model
371 imperfection on CMP quality. The first of these features is tuning. Tuning is the
372 modification of parameterization scheme parameters so as to accommodate – create
373 agreement with – old data. A prominent instance is the setting of parameters
374 associated with the small-scale mixing processes in the ocean. Tuning to current day
375 conditions is hard to avoid given the limited available data about the climate system.
376 Moreover, climate scientists worry that when model success results from
377 accommodation, it provides less confirmation of model abilities than success that
378 results from out-of-sample prediction, that is from prediction that is made prior to the
379 availability of the data but that nevertheless accurately captures the data (Knutti,
380 2008; Smith, 2006; Stainforth et al., 2007a). Prominently, there is the suspicion that
381 accommodation threatens to guarantee success irrespective of whether models
382 correctly capture those underlying processes within the climate system that are
383 relevant to its long term evolution (Schwartz et al., 2007). This impacts assessing
384 projection quality. Difficulty in assessing the extent to which a model's basic
385 assumptions hold will give rise to difficulty in assessing its projections.

386 Work in the philosophy of science, however, shows that whether, and under
387 what conditions, the accommodation of data provides reduced confirmation is an
388 unresolved one (Barrett and Stanford, 2006). On the one hand, some philosophers do
389 worry that accommodation raises the threat of generating empirical success
390 irrespective of whether one's theoretical assumptions are correct (Worrall, 2010). On
391 the other hand, if we prioritize out-of-sample prediction over accommodation,
392 evidence might be good evidence of the suitability of model *A* for generating a set of
393 projections *R* for the late 21st century and not so good evidence for the suitability of
394 model *B* for this purpose even though the models are intrinsically identical. This

395 might occur because the developers of model *B* happen to learn, while those of *A* do
396 not learn, of relevant evidence at the stage of model development. In such
397 circumstances, the developers of *B* might end up accommodating the evidence while
398 the developers of *A* successfully predict it. Resulting differing degrees of confidence
399 in the models would, paradoxically, have to be maintained even if it were recognized
400 that the models are intrinsically identical. If accommodated evidence as such is poor
401 evidence, what determines whether evidence is good evidence for a model is the
402 model's history and not just its intrinsic characteristics (see, e.g., Hudson (2007) for
403 worries about the value of out-of-sample prediction).

404 Unfortunately, while the philosophy of science literature tells us that tuning
405 might not be so bad, it still leaves open the possibility that it is problematic. So how
406 tuning affects CMP accuracy still needs to be addressed.

407 Of course, different approaches to parameterization affect CMP quality
408 differently. For example, stochastic parameterizations, i.e., parameterizations that
409 introduce small but random variations in certain model parameters or variables, are
410 arguably sometimes better than standard deterministic parameterizations (Palmer et
411 al., 2005). The worries about tuning, however, arise for all available parameterization
412 techniques.

413

414 **3.3 The long term nature of projections and its impact on assessing projections**

415 A second factor that, according to some climate scientists, complicates determining
416 the impact of model imperfection is the fact that climate models cannot be tested
417 repeatedly across relevant temporal domains (Frame et al., 2007; Knutti, 2008). We
418 can repeatedly compare weather model forecasts with observations. Success
419 frequencies can then be used to provide probabilistic estimates of model fitness for the

420 purpose of generating accurate forecasts. Recently, some old CMPs have been directly
421 assessed (Hargreaves, 2010). But many CMPs have fulfillment conditions that are
422 never realized and, anyway, CMPs are generally too long term to allow repeated
423 direct testing. Thus, it has been argued, it is hard to take the impact of many model
424 implemented assumptions about long term climate into account in assessing model
425 suitability for generating projections.

426 But the fact that we cannot test our models' predictions over the time scales of
427 the predictions is not itself a difficulty. Consider predictions of Earth orbit variation
428 induced changes in solar radiation at the top of atmosphere over the next million
429 years. Here, predictions are generated using model implemented theory about orbital
430 physics, including Newtonian mechanics and an understanding of its limitations
431 (Laskar et al., 2004). This theory is what grounds confidence in the predictions,
432 though the theory and the models based upon it are only tested against relatively
433 short-term data. As the general views we will discuss about how scientific claims are
434 assessed illustrate, there is no need to assume that estimates of a model's ability must
435 be, or are, made on the basis of numerous observations of how well the model has
436 done in the past.

437

438 **3.4 Basic theory, recognized model imperfection and assessing projections**

439 There are nevertheless two more factors other than tuning that complicate taking into
440 account the effects of model imperfection in assessing projection quality. The first,
441 which is not explicitly discussed in the climate science literature but which climate
442 scientists no doubt recognize, is the combination of known model imperfection with
443 the fact that the background knowledge used in constructing models provides a
444 limited constraint on model construction.

445 Philosophers of science observe that theory provides essential information
446 about model reliability (Humphreys, 2004). Newtonian physics, general relativity and
447 other theories provide essential information about when, and to what extent, we can
448 neglect aspects of the solar system in applying Newtonian theory to model the orbit of
449 the Earth. The same, we have noted, is true of models of how changes in the Earth's
450 orbit affect top of the atmosphere solar radiation. In the case of climate modeling,
451 however, the extent to which theory can guide climate model construction and
452 projection quality assessment is limited. After all, parameterization is introduced
453 precisely because of a limited ability to apply explicit theory in model construction.

454 We do not, for example, have a quantitative theory of the main mechanisms of
455 the stratospheric circulation. As a result, while our partial understanding of these
456 mechanisms can be used in arguing that CMIP3 GCMs' limited ability to represent
457 the stratosphere adversely affects their simulations of tropospheric climate change, the
458 way and extent to which it does so will remain a matter of ongoing investigation (as
459 in, e.g., Dall' Amico (2010)).

460 A limited ability to apply theory in model construction will even make it
461 difficult to decide what we can learn about CMP accuracy from whatever success
462 models have. For easy, relatively theory neutral, ways of drawing conclusions from
463 model successes are hard to come by given model imperfection.

464 Model imperfection implies that models will only have limited empirical
465 success, as indeed is found in the case of climate models. The strongest claim reported
466 by WG1 AR4 on behalf of simulated GCM multi-model annual mean surface
467 temperatures is that, outside of data poor regions such as the polar regions, simulated
468 temperatures were usually within 2 °C of observed temperatures. For most latitudes,
469 the error in simulated zonally averaged outgoing shortwave radiation was about 6%.

470 Simulation of the strength of the Atlantic Meridional Overturning Circulation (MOC)
471 suffers from substantial inaccuracies (Fig. 3). And the same is true of simulation of
472 precipitation patterns, especially on regional scales (Randall et al., 2007). Such
473 inaccuracies short-circuit a simple argument for assigning a high quality to CMPs,
474 namely one that assigns them such a quality on the ground that they were generated
475 by models which simulate data well across the board. Indeed, there is reason to think
476 that increased ability to simulate the current mean climate state across large sets of
477 climate variables is a limited constraint on CMP accuracy (Abe et al., 2009; Knutti et
478 al., 2010). For example, it has been shown (Knutti et al., 2010) that the range of
479 CMPs of precipitation trends is not substantially affected by whether it is produced by
480 all the CMIP3 models or by a subset of high performing models. Assessment of a
481 projection's quality requires correctly identifying which, if any, aspects of model
482 performance are relevant to the projection's accuracy.

483 Further difficulty in figuring out what to infer from what model success there
484 is arises from the well recognized interdependency of climatic processes. Changes in
485 some climatic processes inevitably give rise to changes in others. Changes in cloud
486 cover, land usage, soil hydrology, boundary layer structure and aerosols will, for
487 example, affect surface temperature trends and vice versa. Thus, an accurate
488 simulation of some quantity x will require an appropriate simulation of related
489 quantities upon which x depends. And our assessment of the quality of a projection of
490 x will have to take into account both the accuracy with which x has been simulated
491 and the accuracy with which related quantities have been simulated. One cannot
492 simply argue that since some models simulate a certain climatic quantity well, their
493 projections of this quantity are good (Parker, 2009).

494 Easy, relatively theory neutral ways of assessing what to infer from limited
495 model successes might also be hampered by structural instability, which is, like high
496 sensitivity to changes in initial conditions, a feature of nonlinear systems. A system is
497 structurally unstable when slight changes to its underlying dynamics would give rise
498 to qualitatively different system evolutions. Components of the climate system do
499 exhibit structural instability (Ghil et al., 2008; McWilliams, 2007). This means that
500 minor observed errors in simulating current climate might, given model imperfection,
501 lead to substantial errors in CMPs.

502

503 **3.5 Unrecognized model imperfection and assessing projections**

504 The final source of difficulty for assessing projection quality in light of model
505 imperfection is the possibility, worried about by scientists from all fields, that our
506 models are wrong in unrecognized ways. Empirically successful theories and models
507 have often turned out to rest on mistaken assumptions about which theoretical – that is
508 not directly observable – processes and entities explain observable phenomena
509 (Laudan, 1981). This is true of theories and models of the climate system. Prior to the
510 1990s, for example, climate models that were used to provide spatial simulations of
511 global surface temperatures did not include a representation of the role of aerosols in
512 the climate system and this turned out to be a surprisingly substantial incompleteness
513 in the simulations (Wigley, 1994). Moreover, current candidates for substantially
514 underestimated forcing, feedbacks and internal variability exist (e.g., terrestrial
515 biogeochemical feedbacks (Arneth et al., 2010) and feedbacks amplifying the effects
516 of solar luminosity (Kirkby, 2007)).

517 Some philosophers have concluded, largely on the basis of the history of
518 successful but superseded theories and models, that a theory or model's predictive

519 success should not be used to justify belief in what the theory or model tells us about
520 theoretical entities and processes (see, e.g., Stanford (2006)). On their view, theories
521 and models should be taken to be no more than tools for predicting observable
522 phenomena. The sad truth, however, is that it is currently unclear what we are entitled
523 to assume about how complete empirically successful theories and models are (see
524 Saatsi (2005) and Psillos (1999) for two of many further alternative perspectives on
525 this unresolved issue). In particular, it is unclear what we are entitled to assume about
526 how complete climate models and our knowledge of the climate system are, including
527 about how complete our knowledge of climatic factors that are materially relevant to
528 CMP accuracy is. This complicates assessment. For example, difficulty in estimating
529 the completeness of GCMs' representations of the effects of solar luminosity
530 fluctuations means difficulty in assessing projections of GMST trends.

531

532 **3.6 Criteria of adequacy for approaches to assessing projections**

533 Our discussion of why assessing projection quality is difficult helps to spell out
534 criteria of adequacy for approaches to such assessment. Adequate approaches will,
535 given initial condition inaccuracy, have to assess projection quality in light of the
536 possible path dependent nature of error propagation. Given the inevitable use of
537 parameterization, they will have to take the possible effects of tuning into account.
538 They will also have to take the impact of model imperfection into account. Doing so
539 involves paying attention to climate models' limited ability to simulate climate, to the
540 difficulty in determining which aspects of model empirical success are relevant to
541 assessing which projections, to the interdependence of the evolution of climatic
542 quantities along with the effect of this interdependence on error propagation and to
543 possible structural instability. Doing so also requires attending to the history induced

544 lack of clarity about unrecognized model imperfection. If the claim is that we are
545 entitled to ignore the history of successful but superseded models and thus to cease
546 worrying about unrecognized model imperfection, we need to be told why. Otherwise,
547 the impact of unrecognized climate model limitations on the accuracy of their
548 projections needs to be taken into account.

549 Since we know that only some of the projections of climate models will be
550 accurate, an adequate approach to assessing projection quality will have to provide
551 projection (or class of projections) specific assessments (Gleckler et al., 2008; Parker,
552 2009). It should judge the quality of a CMP on the basis of how fit the model or
553 models which generated it are for the purpose of doing so, i.e., for the purpose of
554 correctly answering the question the CMP answers.

555

556 **4. The confidence building approach**

557 We now discuss the confidence building approach to assessing projection quality.
558 This approach, recall, focuses on what model agreement with physical theory as well
559 as model simulation accuracy confirm. Better grounding in physical theory and
560 increased accuracy in simulation of observed and past climate is used to increase
561 confidence in models and hence in CMPs. Given the emphasis on grounding in
562 physical theory, the reliance here is primarily on GCMs.

563 In the uncertainty assessment guidance note for WG1 AR4 lead authors,
564 degrees of confidence in models are interpreted probabilistically. Specifically, they
565 are calibrated to chance ranges, e.g., very high confidence in a model is interpreted as
566 its having an at least 9 in 10 chance of being correct (IPCC 2005). The chance that a
567 model is correct can be thought of as the model's propensity to yield correct results
568 with a certain frequency, but neither the guidance note nor the report itself indicate

569 how chances should be interpreted. Indeed, they do not indicate how the talk of
570 chances of models' being correct relates to the talk of CMP likelihoods, and the report
571 does not go beyond establishing increased confidence in models in order to assign
572 them specific degrees of confidence. This last fact makes it unclear how the report's
573 use of 'increased confidence' relates to the explication of degrees of confidence in
574 terms of chances. Better grounding in physical theory is illustrated by the, at least
575 partly theoretically motivated, inclusion in some GCMs of interactive aerosol modules
576 (Randall et al., 2007). Illustrations of improved simulation accuracy are given below.

577

578 **4.1 Initial condition inaccuracy and the confidence building approach**

579 WG1 AR4 states that many climatic quantities of interest, including those relating to
580 anthropogenic climate change, are much less prone to nonlinear sensitivity to initial
581 conditions than weather related quantities and are thus more amenable to prediction
582 (Le Treut et al., 2007). This relative insensitivity to initial conditions is argued for
583 primarily on the basis of GCM simulations in which initial conditions are varied.
584 Notably, CMIP3 multi-model simulations of 20th century GMST, in which ranges
585 reflect different initial condition runs of participating models, suggest little internal
586 variability in GMST over periods of decades and almost none over the whole century
587 (See Fig. 1 and (Hawkins and Sutton, 2009)).

588 WG1 AR4 acknowledges that confidence in simulations of response to
589 changes in initial conditions depends on resolving worries about the effects of
590 relevant model imperfection (Meehl et al., 2007b). But the claim is that these worries
591 can be mitigated by examining how well GCMs simulate important sources of the
592 climate system's nonlinear responses, e.g., the El Niño – Southern Oscillation (ENSO)
593 and the MOC. Thus, the ability of GCMs to simulate observed nonlinear change in the

594 Atlantic MOC in response to fresh water influx has been used to argue that they can
595 produce reliable projections of aspects of 21st century MOC behavior but that
596 confidence in projections beyond the 21st century is very limited (Pitman and Stouffer,
597 2006).

598 Computational resources, however, only allowed a very limited range of initial
599 conditions to be explored by CMIP3 GCMs (CMIP3, 2007). As to the question of the
600 extent to which GCM ability to simulate (in)sensitivity to initial conditions does help
601 with assessment in light of model imperfection and tuning, it is addressed in the
602 following sections. Here we only note that the need to address this question has been
603 made pressing since WG1 AR4. Recent work suggests that GCMs do not adequately
604 capture the structure of the climate system prior to abrupt changes in the past and are,
605 in some circumstances, insufficiently sensitive to initial conditions. They can, for
606 example, only simulate the cessation of the MOC under about 10 times of the best
607 estimate of actual fresh water influx that has brought it about in the past (Valdes,
608 2011). There is, in addition, a spate of studies according to which CMIP3 GCMs
609 substantially underestimate the extent to which 20th century GMST anomalies are due
610 to internal variability, including initial condition variability, on multidecadal scales
611 (Semenov et al., 2010; Swanson et al., 2009; Wu et al., 2011). Some work suggests
612 that the underestimates extend to periods of 50 to 80 years in length (Wyatt et al.,
613 2011).

614 Recognizing the potential significance of initial conditions to improving
615 multidecadal CMPs, some recent work aims to take on the challenge of limited
616 available data in order to initialize simulation runs to actual observed initial
617 conditions (Hurrell et al., 2009). More extensive exploration of the impact of varying

618 GCM simulation initial condition settings is also being carried out (Branstator and
619 Teng, 2010).

620

621 **4.2 Parameterization, tuning and the confidence building approach**

622 WG1 AR4 addresses the difficulty of assessing projection quality in light of tuning by
623 taking increased simulation accuracy to increase confidence in models only when this
624 accuracy is not a result of direct tuning, i.e., only when it is not the result of tuning a
625 parameter for a certain quantity to observations of that quantity (Randall et al., 2007,
626 p. 596). But tuning can be indirect. GCMs do not possess parameters for GMST
627 trends, and thus cannot be directly tuned to observations of these trends. Nevertheless,
628 there is (CCSP, 2009) substantial uncertainty about radiative forcings, and especially
629 about aerosol forcing, allowing forcing parameters to be tuned to yield close
630 agreement between simulated and observed 20th century mean GMST trends (Fig. 1).
631 That this tuning occurs is, as is widely recognized within the climate science
632 community, suggested by the observation that different models achieve such
633 agreement by substantially different combinations of estimates of climate sensitivity
634 and radiative forcing [CCSP, 2009; Knutti, 2008b].

635 The difficulty in assessing projection quality in light of parameterization
636 limitations is partly, if implicitly, addressed by noting improvements in
637 parameterization schemes since the publication of WG1 TAR. As schemes that
638 incorporate a better understanding of the climate system and show better agreement
639 with data become available, we acquire a better understanding of the limitations of
640 older schemes and increase trust in model performance. Such improvement, however,
641 leaves open the question of how to handle worries about tuning. Moreover, increased
642 quality of parameterizations does not indicate how to assess the impact of the

643 inevitable remaining underdetermination in parameterization choice on projection
644 quality. Thus, it remains unclear how accurate CMPs actually are.

645 Another strategy that is not explicitly discussed in WG1 AR4, but which is
646 consistent with the confidence building approach, is suggested by the idea that
647 grounding in basic theory increases confidence in models. Perhaps, in some cases, the
648 role of basic theory in generating CMPs is sufficient so as to eliminate, or
649 substantially reduce, worries arising from the use of parameterizations. It has been
650 argued that while simulating the feedback effect of increased water vapor inevitably
651 makes use of parameterizations, this effect is dominated by processes that are
652 represented by the equations of fluid dynamics and thus will continue to be accurately
653 simulated by climate models (Dessler and Sherwood, 2009). It has also been
654 suggested that, since GCMs use the equations of fluid dynamics, our ability to predict
655 nonlinear MOC evolution that results from its fundamental properties is beginning to
656 mature, unlike our ability to predict nonlinear evolution it might exhibit as a result of
657 terrestrial ecosystems (Pitman and Stouffer, 2006).

658 One difficulty here is how to determine that properties represented by basic
659 physical theory largely determine the evolution of projected quantities. Insofar as
660 estimates that this is so rely on – as, e.g., Dessler and Sherwood (2009) rely on –
661 climate model results, it is assumed that available parameterizations are adequate and
662 the reliance on parameterization is not bypassed. Further, even if we have managed to
663 isolate properties that are represented by basic theory and determine the evolution of a
664 projected quantity, we cannot escape worries relating to the use of parameterization.
665 Parameterization always plays an essential role even in descriptions of subsystems of
666 the climate for which we possess basic equations. Basic equation discretization in
667 GCMs brings with it grid-scale dependent parameterization, e.g., grid-scale dependent

668 convection parameterization, of subgrid processes. How this discretization and
669 associated parameterization affects CMP accuracy, especially in light of how it affects
670 model ability to simulate highly nonlinear dynamics, needs adequate treatment.

671

672 **4.3 Structural inadequacy and the confidence building approach**

673 Increased model grounding in basic physical theory and increased accuracy in
674 simulation results across a range of such results does indicate increased structural
675 adequacy. Moreover, confidence building exercises do typically acknowledge a wide
676 variety of model limitations. What we need, however, are arguments connecting
677 increased success with the quality of specific classes of CMPs. This includes
678 arguments addressing the issue of how total remaining inadequacy affects CMP
679 quality.

680 Thus, for example, WG1 AR4 offers information such as that more state-of-
681 the-art models no longer use flux adjustments, that resolution in the best models is
682 improving, that more physical processes are now represented in models and that more
683 such processes are explicitly represented (Randall et al., 2007). But we need
684 arguments that connect these successes to an overall estimate of remaining structural
685 inadequacy and tell us what this inadequacy means for the quality of specific classes
686 of CMPs. It is one thing to be shown that simulated multi-model mean surface
687 temperatures are, outside of data poor regions, usually within 2 °C of observed
688 temperatures, another to be shown how this information bears on the quality of CMPs
689 of mean surface temperature trends and yet another to be shown how it bears on the
690 quality CMPs of mean precipitation trends.

691 While the needed arguments can be further developed, it remains to be seen
692 how far they can be developed. Further, it is likely that these arguments will, to a

693 substantial extent, be based on theory and expert judgment, thus limiting the extent to
694 which the confidence building approach is model based.

695

696 **4.4 The appeal to paleoclimate**

697 An important distinction needs to be made between model ability to simulate 20th
698 century climate and model ability to simulate paleoclimate. The latter provides
699 opportunities for out-of-sample testing, as WG1 AR4 notes (Jansen et al., 2007, p.
700 440). Such testing is of particular significance as it has the potential to help in
701 addressing the question of the extent to which tuning to current climate is a problem.
702 Indeed, there is growing recognition of the importance of palaeodata, including of its
703 importance for model assessment (Caseldine et al., 2010). In this context, there is an
704 ongoing debate about whether to conclude that GCMs lack representations of crucial
705 mechanisms/feedbacks because these models have difficulties in accurately
706 simulating past warm, equable climates with a weak equator-to-pole temperature
707 gradient (Huber and Caballero, 2011; Spicer et al., 2008).

708 Although this may change in the future, the burden of assessing models in
709 light of data nevertheless currently rests firmly on the ability of models to simulate
710 recent climate. This is so for at least three reasons. First, simulation experiments with
711 paleodata are still limited. WG1 AR4's appeal to such simulations is confined
712 primarily to two instances. WG1 AR4 uses model ability to simulate aspects of the
713 climate system during the Last Glacial Maximum (LGM) in order further to support
714 the claim that models have captured the primary feedbacks operating in the climate
715 system at the time (Jansen et al., 2007, p. 452). WG1 AR4 also uses model ability to
716 simulate climate responses to orbital forcing during the mid-Holocene in order to
717 improve confidence in model ability to simulate responses to such forcing (Jansen et

718 al., 2007, p. 459). Second, most of the models WG1 AR4 relies on in generating
719 projections are not among the models it relies on in discussing paleoclimate
720 simulations (Schmidt, 2010). And when the same models are relied on in both
721 contexts, model resolution usually varies across the contexts (Braconnot et al., 2007).
722 Practical constraints mean lower resolution models have to be used to simulate
723 paleoclimate. Thus it is unclear what the paleoclimate simulation successes allow us
724 to conclude about model fitness for the purpose of generating projections. Third, there
725 are substantial, unresolved issues about how uncertain paleoclimate reconstructions
726 are, and thus about what we can learn from them (Snyder, 2010; Wunsch, 2010).

727

728 **4.5 Inter-model results, robust projections and the confidence building approach**

729 The confidence building approach is strengthened, both in WG1 AR4 and elsewhere,
730 by noting that state-of-the-art GCMs provide a robust and unambiguous picture of the
731 evolution of some large scale features of climate. Such multi-model results are
732 supposed to increase confidence in projections. For example, state-of-the-art GCMs
733 predict that GMST evolution will be roughly linear over much of this century, thus
734 supposedly reducing worries about the sensitivity of such evolution to initial condition
735 changes and to minor variations in model structure (Knutti, 2008).

736 How does the appeal to multi-model results help in assessing projection
737 quality, as opposed to improving projection accuracy? We outline two views about
738 how it does so and then critically discuss these views.

739 A common assumption in formal analyses of multi-model ensemble results,
740 and to some extent in applications of the confidence building approach, is that model
741 errors are independent of each other and thus tend to cancel out in calculations of
742 multi-model means (Meehl et al., 2007b; Palmer et al., 2005; Tebaldi and Knutti,

743 2007). Indeed, there is empirical evidence that multi-model means are more accurate
744 than are the results of individual models (see Gleckler et al. (2008) as well as, for
745 further references, Knutti et al. (2010)). Given the assumptions of error independence
746 and of error cancellation, one could argue that we can expect a reduction of error in
747 ensemble means with increased model numbers and thus can take the number of
748 models used in generating means to be an indicator of CMP quality (Tebaldi and
749 Knutti, 2007).

750 In addition, or alternatively, one can assume that ensemble models are to some
751 extent independent of each other in that they explore alternative model structures and
752 parameterizations that are consistent with our knowledge of the climate system
753 (Murphy et al., 2007). Ensemble projection ranges can then be viewed as at least
754 partial explorations of our uncertainty about the climate system and can thus be used
755 to tell us something about projection quality. One might suggest, in particular, that the
756 greater the extent to which the range of uncertainty is explored by an ensemble, the
757 greater the extent to which the projections/projection ranges it produces are robust or
758 insensitive to uncertain assumptions and thus the more probable these results are
759 (Weisberg (2006) describes the general logic behind appeals to robustness). Multi-
760 model ensemble projection ranges are sometimes interpreted probabilistically, e.g.,
761 the range of generated projections is supposed to span the range of possibilities and
762 each projection is assigned a probability equal to the fraction of models that generate
763 it (as in Räisänen and Palmer (2001) and, to some extent, in WG1 TAR (IPCC 2001)).

764 The appeal to multi-model results does not, and is not intended to, address the
765 issue of tuning or the difficulty of figuring out what to infer about the quality of
766 specific CMPs from the partial empirical successes of models. Further, worries about

767 the use of multi-model ensembles have been raised both within and without climate
768 science.

769 Philosophers have pointed out that individual model error can only cancel out
770 to a limited extent because limited knowledge and limited computational resources
771 mean that where one model's error is not repeated by another model, the other model
772 will probably have to introduce a different error (Odenbaugh and Alexandrova, 2011).
773 Limited knowledge and limited computational resources also mean that substantial
774 model imperfection will inevitably be shared across models in ensembles (Odenbaugh
775 and Alexandrova, 2011). Multi-model ensembles in all fields of research accordingly
776 inevitably leave us with substantial error the impact of which on results is not
777 estimated. So, while coming to rely on multi-model ensembles might entitle us to be
778 more confident in projections than we would have been otherwise, it does not appear
779 to allow us to assign qualities that, like probabilities and informative probability
780 ranges, involve specifying the full range of possible evolutions of projected quantities.

781 Climate scientists' examination of GCM ensemble results confirms that such
782 ensembles only provide limited improvement in agreement with empirical data and
783 that much of the remaining disagreement arises from biases that are systematic across
784 ensemble members (Knutti et al., 2010). For present day temperature, for example,
785 half of the bias exhibited by the ensemble of models used by CMIP3 would remain
786 even if the ensemble were enlarged to include an indefinite number of models of
787 similar quality (Fig. 4). The observation that models share model imperfections is also
788 acknowledged in climate science research, including in WG1 AR4. Climate modelers
789 tend to aim at constructing the best models they can for their shared purposes and in
790 doing so inevitably use shared knowledge and similar technology. As a result, climate
791 models tend to be similar, sharing many of the same imperfections (Allen and Ingram,

792 2002; Knutti, 2010; Meehl et al., 2007b; Stainforth et al., 2007a; Tebaldi and Knutti,
793 2007).

794 A related problem is that, although model limitations are extensively examined
795 in the literature, discussion of the extent to which models in specific multi-model
796 ensembles differ in ways that are relevant to assessing projections is limited (Knutti et
797 al., 2010).

798 Recognizing the limited extent to which model error cancels out, some climate
799 scientists have suggested that we should not assume that the larger the ensemble the
800 closer means are to representing reality. Instead, they suggest, one should assume that
801 the correct climate and the climates simulated by models in an ensemble are drawn
802 from the same distribution, e.g., from the standard normal (Gaussian) distribution.
803 Under this new assumption, the failure of an increase in ensemble size to improve
804 simulation results is no longer interpreted as indicating systematic bias. One can then,
805 the suggestion is, assume that when a proportion r of an ensemble yield a given
806 projection, r is the probability of that projection (Annan and Hargreaves, 2010). But
807 the assumption that model probability distributions coincide with the real climate
808 distribution cannot be made in general, as is illustrated in the case of the already
809 mentioned GCM inability realistically to simulate historical Atlantic MOC collapse.
810 Indeed, structural inadequacy that is known to be shared by ensemble models means
811 that we know that the correct climate *cannot* be represented by current models.

812 Let us now look at the second argument for appealing to inter-model results in
813 assessing projection quality, the one according to which multi-model ensembles allow
814 us to explore our uncertainty. Since existing climate models share many uncertain
815 assumptions, the projections/projection ranges multi-model ensembles produce do not
816 reflect full explorations of our uncertainty (Parker, 2011; Pirtle et al., 2010).

817 Moreover, once again, such ensembles do not allow assigning projection qualities the
818 assignment of which involves estimating the full range of possible evolutions of
819 projected quantities.

820 The GCMs used by WG1 AR4 only sample some of the recognized range of
821 uncertainty about aerosol forcing, perhaps because of the already mentioned tuning
822 relating to this forcing. As a result, the spread of estimated temperature anomalies
823 these models provide (Fig. 1) substantially underestimates the uncertainty about this
824 anomaly and, accordingly, would be misleading as a guide to projection quality
825 (Schwartz et al., 2007). So too, if we take the range of natural variability covered by
826 the simulations represented in Fig. 1 to reflect our uncertainty about natural variability
827 over the next three decades, we will assign a very low probability to the prediction
828 that natural variability will substantially affect GMST trends over this period.
829 Keeping in mind, however, that these models may well similarly and substantially
830 underestimate internal variability over the next 30 years would lead us to reduce our
831 confidence in this prediction. Worse, if we cannot estimate the probability that the
832 ensemble is wrong (something the ensemble cannot help us with!) about internal
833 variability here, we are not in a position to assign the prediction a probability.

834 A number of suggestions have been made within the climate science
835 community about how partially to address the above worries about the use of multi-
836 model ensembles. Assessments that are explicit about the extent to which climate
837 models in any multi-model ensemble differ in ways that are relevant to assessing
838 projection quality should be offered (IPCC 2010; Knutti et al., 2010). If, for example,
839 internal variability in the MOC is an important source of uncertainty for projections of
840 mean sea surface temperatures over the next 30 years and our ensemble is in the
841 business of making such projections, it should be clear to what extent the simulations

842 produced by the ensemble differ from each other in ways that explore how internal
843 variability in the MOC might occur. Assessing projection quality relevant differences
844 in models is a substantial task, one that goes well beyond the standard multi-model
845 exercise.

846 In addition, while limited knowledge and resources, e.g., restrictions to certain
847 grid resolutions, mean that there is no question of exploring all of existing uncertainty,
848 provision of second and third best guess modeling attempts could provide a clearer
849 picture of our uncertainty and its impact on CMP quality (Knutti et al., 2010; Smith,
850 2006).

851 A difficulty to keep in mind is that of determining how a model component
852 that is shared by complex models that differ in complex ways affects CMP quality.
853 Assessment of model components and their impact on model performance is a
854 challenge that is – because of the need to evaluate models in light of background
855 knowledge – part and parcel of assessing models fitness for purpose. This challenge is
856 complicated when the projection is generated by complex models that implement
857 common components but differ in other complex ways. For the same component may,
858 as a result, function in different ways in different models (Lenhard and Winsberg,
859 2010). Examining how a parameterization of cloud microphysics affects CMPs may,
860 for example, be hampered if the parameterization scheme is embedded in models that
861 substantially differ in other parameterizations and/or basic theory.

862 The comparison of substantially differing models will also exacerbate existing
863 challenges for synthesizing the results of multi-model ensembles. Climate scientists
864 have noted that synthesizing the results of different models using a multi-model mean
865 can be misleading even when, as in the case of the CMIP3 models, the models
866 incorporate only, and only standard, representations of atmosphere, ocean, sea ice and

867 land [Knutti et al., 2010]. For example, the CMIP3 multi-model mean of projected
868 local precipitation changes over the next century is 50 % smaller than that which
869 would be expected if we were to assume that at least one, we know not which, of the
870 CMIP3 models is correct. So it seems that using a mean in this case is misleading
871 about what the models describe (Knutti et al., 2010). Synthesizing the results of
872 different models may be even more misleading where models differ substantially in
873 how they represent processes or in which processes they represent, e.g., if some of the
874 models do and some do not include representations of biogeochemical cycles (Tebaldi
875 and Knutti, 2007). In such circumstances, for example, a mean produced by two
876 models may well be a state that is impossible according to both models.

877

878 **5. The subjective Bayesian approach**

879 Perhaps the main approach to supplement the confidence building approach in WG1
880 AR4 is the subjective Bayesian approach. We first consider this formal,
881 supplementary approach as it is used to assess projection quality in light of difficulties
882 in parameter choice (Hegerl et al., 2006; Murphy et al., 2004). We then consider how
883 it has been extended.

884

885 **5.1 The subjective Bayesian approach to parameter estimation**

886 A simple, but representative, application of the standard version of the Bayesian
887 approach to parameter, including projection parameter, estimation involves
888 calculating the posterior probability distribution function $P(F | \text{data}, M)$ using Bayes'
889 theorem, as in Eq. (3) (Frame et al., 2007). $P(F | \text{data}, M)$ specifies the probabilities
890 of values of a parameter, F , given data and a model M . $P(\text{data} | F, M)$ is the likelihood
891 of F and captures, as a function of values of F , the probability that the data would be

892 simulated by M . In the Bayesian context, ‘the likelihood of F ’ refers to a probability
 893 function for data rather than, as it would on the WG1 AR4 use of ‘likelihood’, to a
 894 probability range for F . The prior probability distribution function $P(F | M)$ is the
 895 probability distribution function of F given only M and thus prior to consideration of
 896 the data. $P(\text{data})$ is a normalizing constant required to ensure that the probabilities
 897 sum up to 1.

$$898 \quad P(F | \text{data}, M) = P(\text{data} | F, M)P(F | M)/P(\text{data}) \quad (3)$$

899

900 The probabilities in Eq. (3) are, on the subjective Bayesian approach, to be
 901 interpreted as precise, quantitative measures of strength of belief, so called ‘degrees of
 902 belief’. What makes the subjective Bayesian approach subjective is that unconstrained
 903 expert opinion – the beliefs of certain subjects irrespective of whether they meet
 904 objective criteria of rationality such as being well grounded in empirical evidence – is
 905 used as a central source for selecting prior probability distributions. Still, the
 906 subjective Bayesian approach often uses uniform assignments of priors. In doing so, it
 907 borrows from what is usually called ‘objective Bayesianism’ (see Strevens (2006b) for
 908 a discussion of the different forms of Bayesian approaches to science).

909 Bayes’ theorem allows us to take existing estimates of parameter uncertainty –
 910 here captured by $P(F | M)$ – and to constrain these using information from perturbed
 911 physics experiments about how well a model simulates data as a function of parameter
 912 settings – information here captured by the likelihood function $P(\text{data} | F, M)$.
 913 Assume experts provide prior probability distributions for parameters relating to total
 914 radiative and present-day indirect aerosol forcing and that we calculate the probability
 915 that a model gives, as a function of the parameters’ values, to observed oceanic and
 916 atmospheric temperature change. Bayes’ rule can then yield posterior probability

917 distributions for the parameters (Fig. 5). Bayesian parameter estimation has tended to
918 rely on models of intermediate complexity and on energy balance models.

919 The Bayesian hope is that the constraints provided by simulation success on
920 parameter estimates will increase the objectivity of such estimates. Moreover, Bayes'
921 theorem provides, what the confidence building approach does not provide, a clear
922 mechanism that relates simulation accuracy to conclusions about CMP quality, thus
923 helping to address the problem of what to infer from available simulation accuracy
924 given the existence of model imperfection.

925 Nevertheless, the standard version of the Bayesian approach to parameter
926 estimation faces substantial problems. The standard interpretation of the probability
927 distributions $P(F | M)$ and $P(F | \text{data}, M)$ is that they are probability distributions for F
928 that are conditional on the correctness of a version of M . In the present context, what
929 is being assumed to be correct is a model version in which one or more parameters are
930 unspecified within a certain range. For the goal is to select parameter values from
931 within a range of such values. Now, it is on the basis of the standard interpretation of
932 $P(F | M)$ and $P(F | \text{data}, M)$ that standard justifications, using so-called Dutch Book
933 arguments, for updating beliefs in accord with Bayes' theorem proceed. Dutch Book
934 arguments generally assume that the, typically statistical, model versions upon which
935 probabilities are conditional are correct. It is argued that, given this assumption, the
936 believer would end up with beliefs that are not as true as they might have been, or
937 would incur a financial loss, if his or her beliefs were not updated in accord with
938 Bayes' theorem (see Jeffrey (1990) and Vineberg (2011) for examples). But if, as in
939 the cases we are concerned with, the model version upon which distributions are
940 conditional is not correct, applying Bayes' theorem may offer no advantage and may
941 be a disadvantage.

942 Assume that our subject relies on a CMIP3 GCM to determine whether a
943 specified fresh water influx will lead to a collapse in the MOC and that the specified
944 influx is a tenth of that needed to get the model to simulate collapse. Assume also that
945 some exploration of plausible parameter settings in the GCM does not alter results
946 substantially. Applying Bayes's theorem on the assumption that the model is, up to
947 plausible parameter modification, correct means that the probability we assign the
948 outcome 'collapse' is 0. The modeler acquiesces to the theorem. Unfortunately, as we
949 now know, the model's results are misleading here. In this case, not applying Bayes'
950 theorem may lead to more realistic judgments.

951 Thus, the standard use of Bayes' theorem in parameter estimation requires an
952 alternative to the standard interpretation of its conditional probabilities. We will also
953 need an alternative to the standard justifications for applying Bayes' theorem.

954 Even if we have settled on some interpretation of the conditional posterior
955 probabilities produced by Eq. (3), there remains the question of what we can infer
956 about reality from these probabilities. There remains, in other words, the question of
957 what distribution of probabilities for F , $P(F)$, we should adopt given the conditional
958 distribution $P(F | \text{data}, M)$. We might have a probability distribution for climate
959 sensitivity that is conditional on the data and a model. But what should we infer from
960 this about actual climate sensitivity? We cannot properly answer such questions until
961 we have gone beyond assessing how parameter choice affects projection quality and
962 have also assessed how structural inadequacy, parameterization scheme choice and
963 initial condition inaccuracy do so (Rougier, 2007).

964 Rougier provides a non-standard version of the Bayesian approach to
965 parameter estimation that has the substantial advantage of allowing us to factor in
966 estimates of structural inadequacy into subjective Bayesian parameter estimates

967 (Rougier, 2007). Nevertheless, his work takes estimates of structural inadequacy as
968 given and thus does not, by itself, tell us how more comprehensive assessments of
969 projection quality are to be produced.

970 Additional difficulties for the Bayesian approach relate to the usage of prior
971 probabilities. We rehearse two familiar worries about this usage. First, estimates of
972 $P(F | M)$ are usually made after data that bears on the estimates is in hand and it is
973 hard to estimate what probability distribution would be assigned to F independently of
974 knowledge of this data. Failure properly to estimate $P(F | M)$ may lead to counting
975 the same data twice, once in estimating priors and once in estimating likelihoods
976 (Frame et al., 2007).

977 Second, while some climate scientists have argued that the explicit setting out
978 of subjective priors by experts is desirable because it makes subjective judgments
979 explicit (Hargreaves, 2010), philosophers of science have pointed out that it leaves
980 open the question of the extent to which experts' views are evidence based and thus
981 puts reliable and unreliable priors on a par (Sober, 2002). This issue becomes
982 particularly worrying in the context of climate modeling. We know that prior selection
983 may be based on results involving tuning and be required even when data
984 underdetermines parameter value choice. So there is a risk that assigning a prior to a
985 parameter value will beg the question against alternative choices and thus yield
986 estimates of climatic variables we are by no means obliged to accept. The worry of
987 question begging is exacerbated by arguments to the effect that the influence of
988 likelihoods, and thus of data, on the shape and width of prior distributions is often
989 minor (Frame et al., 2005).

990 A common way of trying to minimize the impact of the appeal to expert
991 opinion is to represent the state of ignorance that existed prior to the consideration of

992 likelihoods using uniform prior distributions within expert specified ranges. We have
993 already seen that uniform distributions are not suitable for representing ignorance.
994 Moreover, to assume a uniform prior distribution will often be to ignore knowledge
995 we have of the relative plausibility of various prior assignments (Annan and
996 Hargreaves, 2011; Rougier, 2007). So too, a uniform assignment of priors for one
997 parameter will sometimes, because of the non-linear relationship between some model
998 variables, provide a non-uniform prior assignment to another (Frame et al., 2005). It
999 has been suggested that best practice given the worries about prior selection is to
1000 provide readers with posteriors as well as likelihoods. This would somewhat clarify
1001 the role data actually have had in determining posteriors (Frame et al., 2007).

1002 Another way in which the influence of priors might be minimized is by
1003 repeated updating of posteriors in response to new evidence over time. As already
1004 noted, however, evidence with which to test models is mostly limited to familiar 20th
1005 century datasets. There is thus currently limited scope for successive updating of
1006 priors.

1007 As to the idea that the appeal to likelihoods in deriving posterior probabilities
1008 will provide an objective constraint on parameter selection, it also has problems.
1009 Likelihoods measure agreement with data, irrespective of whether such agreement
1010 results from tuning (Katzav, 2011). In addition, we have seen that an adequate
1011 assessment of projection quality needs to take into account not only agreement with
1012 data, but also how error for each simulated quantity develops over projection
1013 scenarios as a function of error associated with other such quantities. Finally, there are
1014 various likelihood metrics or ways of measuring agreement with data. Choice between
1015 these and how such choice affects posteriors is only beginning to be explored (see,
1016 e.g., Ishizaki et al. (2010)).

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1018
1019

5.2 The subjective Bayesian approach and multi-model ensembles

1020 The subjective Bayesian approach has been extended to assessing multi-GCM
1021 ensemble output. This extension, which will be called the subjective Bayesian MM
1022 approach, involves taking an ensemble and producing a statistical model of its
1023 simulation results. Comparing the statistical model and available data yields a
1024 likelihood function that captures the probability the ensemble gives to the data. Bayes'
1025 theorem can then be used, in conjunction with the likelihood function and estimates of
1026 prior probability distributions for the statistical model's parameters, in order to
1027 produce a posterior probability distribution for these parameters (Furrer et al., 2007a;
1028 Furrer et al., 2007b; Leith and Chandler, 2010; Tebaldi et al., 2005; Tebaldi and
1029 Knutti, 2007).

1030 Some variants of the subjective Bayesian MM approach give each ensemble
1031 model equal weight in calculating ensemble posterior probability distributions (Leith
1032 and Chandler, 2010). Other variants weight the contribution of each ensemble model
1033 to posteriors as a function of how well the model simulates aspects of the climate
1034 system (Tebaldi et al., 2005).

1035 Many analyses, e.g., those in WG1 TAR and some of those in WG1 AR4, of
1036 multi-model ensemble results produce projections that are just averages of individual
1037 model results and that have uncertainty ranges which reflect inter-model variability.
1038 This does not yield probabilistic estimates of multi-model ensemble results. The
1039 subjective Bayesian MM approach does yield such estimates. The hope is that doing
1040 so helps to take into account structural inadequacy and limited knowledge of how to
1041 select parameterization schemes. The subjective Bayesian MM approach does not
1042 explicitly tackle the issue of how initial condition inaccuracy affects CMP quality.

1043 The subjective Bayesian MM approach suffers from many of the problems of
1044 the subjective Bayesian approach to parameter estimation. The subjective Bayesian
1045 MM approach faces the problems that arise from the use of prior probabilities. It also
1046 suffers from the problems relating to the choice of likelihood metrics and the failure
1047 to take into account how error for each simulated quantity develops as a function of
1048 error associated with other such quantities. Even weighting models in assessing
1049 projection quality is not a clear advantage given that the data used to do so may well
1050 have already been used in model construction.

1051 Finally, there remain the issues of how to interpret the conditional
1052 probabilities used in Bayes' theorem given model imperfection and of how the
1053 conditional probabilities produced by Bayes' theorem relate to unconditional
1054 probabilities. On the subjective Bayesian MM approach, one updates priors on the
1055 assumption that the statistical model of multi-model ensemble results is correct.
1056 However, given that we know that multi-model ensemble results are biased, this
1057 assumption is false. And any inference from probabilities that are conditional upon
1058 data and an ensemble to unconditional probabilities can only be made given a full
1059 assessment of the effects of initial condition error and model imperfection on CMP
1060 accuracy. We have seen, however, that multi-model ensembles do not provide such an
1061 assessment.

1062

1063 **6. The likelihood approach**

1064 One response to the subjective Bayesian approach's difficulties with subjective prior
1065 probabilities is to try to avoid the use of priors all together. This is what the likelihood
1066 approach does using GCMs. It aims to produce probability distributions for
1067 parameters solely in light of how well models simulate data as a function of parameter

1068 settings, that is solely in light of likelihood functions such as $P(\text{data} | F, M)$ (Allen et
1069 al., 2006). Doing so requires not discounting any parameter settings prior to
1070 simulation and thus providing likelihood functions that span a much broader range of
1071 parameter values than is usual. This has become possible, though usually only in
1072 experiments that perturb the parameters of a single model structure, with the
1073 distributed computing techniques used by climateprediction.net (Frame et al., 2007).
1074 The results of such attempts are distributions that are less biased due to those
1075 parameters that are perturbed, but that are far broader than those otherwise produced.

1076 An application of the likelihood approach is as follows: we take the climate
1077 sensitivities of each of a multi-thousand climateprediction.net ensemble of GCM
1078 variants and estimate the true climate sensitivity to be a weighted sum of these
1079 sensitivities. The weight of each sensitivity is determined by the probability the
1080 variant it belongs to gives to observations of a number of climatic quantities,
1081 including mean sea level temperature, precipitation and surface heat fluxes (Piani et
1082 al., 2005).

1083 The likelihood approach can also be used to minimize the impact of structural
1084 inadequacy and uncertainty about choice of parameterization scheme on CMP
1085 accuracy. It can do so by producing assessments that are only based on the best
1086 simulations available for specific parameter settings (Sanderson et al., 2008). But
1087 focusing on best results does not take into account how they are affected by initial
1088 condition inaccuracy, tuning or aspects of model imperfection other than parameter
1089 choice uncertainty. The same is true of what might be called the multi-model
1090 likelihood approach. This approach uses correlations between GCMs' predictions of
1091 trends for a quantity and related observations formally to select the best predictions
1092 (Boe et al., 2009; Shukla et al., 2006).

1093

1094 **7. Putting it all together**

1095 As we have noted, WG1 AR4 often uses expert judgment that takes the results of the
1096 approaches we have been discussing, as well as partly model-independent approaches,
1097 into consideration in assigning final projection qualities. Insofar as final assignments
1098 are model based, however, the shared limitations of the approaches we have been
1099 discussing remain untouched. In particular, insofar as final assessments are model
1100 based, they face serious challenges when it comes to assessing projection quality in
1101 light of structural inadequacy, tuning and initial condition inaccuracy. Moreover, they
1102 continue to be challenged by the task of assigning probabilities and informative
1103 probability ranges to projections.

1104

1105 **8. Assessing projections: what else can be done?**

1106 We now examine approaches that differ from those that play center stage in WG1
1107 AR4. The first approach, the possibilist approach, is described in the climate science
1108 literature but is primarily non-probabilistic. The remaining approaches are
1109 philosophy-of-science-based approaches. There are currently four main, but not
1110 necessarily mutually exclusive, philosophical approaches to assessing scientific
1111 claims. One of these is the already discussed subjective Bayesian approach. The other
1112 three are those that are discussed below.

1113

1114 **8.1 The possibilist approach**

1115 On the possibilist approach, we should present the range of alternative projections
1116 provided by models as is, insisting that they are no more than possibilities to be taken
1117 into account by researchers and decision makers and that they provide only a lower

1118 bound to the maximal range of uncertainty (Stainforth et al., 2007a; Stainforth et al.,
1119 2007b). Climate model results should, accordingly, be presented using plots of the
1120 actual frequencies with which models have produced specific projections (as in Fig.
1121 6). At the same time, one can supplement projected ranges with informal, though
1122 sometimes probabilistic, assessments of confidence in projections that appeal, as the
1123 confidence building approach appeals, to inter-model agreement and agreement with
1124 physical theory (Stainforth et al., 2007a).

1125 Informal approaches to assessing projection quality must address the same
1126 central challenges that quantitative approaches must address. So, insofar as the
1127 possibilist position allows informal probabilistic assessments of projection quality, it
1128 must address the difficulties that all probabilistic approaches face. However, one
1129 could easily purge the possibilist approach of all probabilistic elements and assess
1130 projections solely in terms of their being possibilities. Moreover, there are obvious
1131 ways to develop purely possibilistic assessment further. Purely possibilistic
1132 assessment can, in particular, be used to rank projections. Possibilities can, for
1133 example, be ranked in terms of how remote they are.

1134 The purged possibilist approach would still face challenges. Presenting CMPs
1135 as possibilities worthy of consideration involves taking a stance on how CMPs relate
1136 to reality. For example, if we are presented with an extreme climate sensitivity range
1137 of 2 to 11 K (Fig. 6) and are told that these are possibilities that should not have been
1138 neglected by AR3 WG1's headline uncertainty ranges (Stainforth et al., 2005), a claim
1139 is implicitly being made about which climate behavior is a real possibility. It is
1140 implied that these possibilities are unlike, for example, the possibility that the United
1141 States will more than halve its budget deficit by 2015. Thus a possibilist assessment of
1142 projection quality needs to be accompanied by an examination of whether the

1143 projections are indeed real possibilities. The same considerations apply to ‘worst case
1144 scenarios’ when these are put forward as worthy of discussion in policy settings or
1145 research. The threat that arises when we do not make sure that possibilities being
1146 considered are real possibilities is that, just as we sometimes underestimate our
1147 certainty, we will sometimes exaggerate our uncertainty.

1148 Nevertheless, the challenges facing purely possibilistic assessment are
1149 substantially more manageable than those facing probabilistic assessment. To say that
1150 something is a real possibility at some time t is, roughly, to say that it is consistent
1151 with the overall way things have been up until t and that nothing known excludes it
1152 (see Deutsch (1990) for a similar definition). A case for a projection’s being a real
1153 possibility can, accordingly, be made just by arguing that we have an understanding of
1154 the overall way relevant aspects of the climate system are, showing that the
1155 projection’s correctness is consistent with this understanding and showing that we do
1156 not know that there is something that ensures that the projection is wrong. There is, as
1157 observed in discussing probabilistic representations of ignorance, no need to specify a
1158 full range of alternatives to the projection here. Further, state-of-the-art GCMs can
1159 sometimes play an important role in establishing that their projections are real
1160 possibilities. State-of-the-art GCMs’ projections of GMST are, for example and given
1161 the extent to which GCMs capture our knowledge of the climate system, real
1162 possibilities.

1163

1164 **8.2 The critical approach**

1165 The first philosophy-of-science-based approach that is not discussed in the IPCC
1166 reports and that will be discussed here is the critical approach (Freedman, 2009;
1167 Longino, 1990). According to this approach, scientific claims are rational to the extent

1168 that they result from open, critical discussion. Longino offers a prominent view of
1169 what such discussion involves. She holds that open critical discussion occurs within a
1170 community to the degree that the community has recognized avenues for criticism of
1171 evidence, methods, assumptions and reasoning; the community's members share
1172 standards of criticism; the community is responsive to criticism and intellectual
1173 authority is shared equally among qualified members (Longino, 1990). Petersen offers
1174 what can be thought of as a version of the critical approach, one that is designed to
1175 assist in, among other things, assessing CMP quality. He provides procedures, and a
1176 classification of types of uncertainty, that are supposed to help systematizing
1177 qualitative assessments of model assumptions and thus to facilitate open, critical
1178 discussion of the quality of model-based-claims (Petersen, 2012).

1179 The motivation for the critical approach is twofold. On the one hand,
1180 according to its proponents, critical discussion allows overcoming individual
1181 subjective bias. On the other hand, there are no available standards beyond our current
1182 standards by which scientific claims can be judged. So, it is argued, rationality cannot
1183 amount to more than the application of available standards of critical discussion and
1184 the acceptance of the deliverances of these standards.

1185 The critical approach is not really an alternative to the approaches used in
1186 WG1 AR4. Rather it is a framework that tells us in what conditions the deliverances
1187 of these approaches are acceptable. Petersen's framework could, for example, be used
1188 to guide applying the confidence building approach.

1189 Further, according to the critical approach, we can recognize that an
1190 assessment of the quality of a projection is limited while nevertheless accepting the
1191 projection. For, on this approach, where acceptance of models' fitness for the purpose
1192 generating projections is a result of open, critical discussion, accepting the models'

1193 projections is reasonable even if the discussion in question has substantial limitations,
1194 e.g., if the impact of unknown structural inadequacy on the projections has not been
1195 taken into account. The critical approach would thus, for example, warrant trust in
1196 state-of-the-art GCMs for the purpose of generating the GMST projections presented
1197 in Fig. 2, subject to expert correction in light of known GCM limitations and provided
1198 that the trust results from open, critical discussion.

1199 Acceptance of models' fitness for purpose can, however and as Longino's
1200 criteria for such criticism state, only be the result of open, critical discussion if there
1201 are shared standards for assessing fitness for purpose. In the absence of shared
1202 standards, agreement will be the result of the arbitrary preference of some standards
1203 over others rather than the uptake and assessment of relevant alternatives. In the case
1204 of CMP assessment, what we need for acceptance of model fitness for purpose to be
1205 the result of open, critical discussion is agreement about issues such as whether
1206 assessment should be probabilistic, whether it should be formal and so on. The present
1207 paper makes it clear, however, that it would be premature to agree on these issues and,
1208 indeed, that there is no such agreement.

1209 A more general worry about the critical approach is that, by itself, it leaves
1210 unaddressed the question of when the results of open, critical discussion are reliable
1211 (Crasnow, 1993). Unless we have an assessment of how reliable current critical
1212 discussion of model fitness for purpose is, it is unclear why we should accept the
1213 results of such discussion.

1214

1215 **8.3 Inference to the best explanation and climate model evaluation**

1216 The next philosophy based approach to assessing projection quality is the inference to
1217 the best explanation (IBE) approach (Lipton, 2004). In discussing the confidence

1218 building approach we saw model confidence being increased on the basis of
1219 improvement in model virtues such as agreement with background knowledge
1220 (including grounding in basic theory), increased realism, agreement with observations
1221 and model scope – that is, roughly, the number of distinct classes of facts a model
1222 simulates. An additional model virtue that is appealed to in climate modeling
1223 (Shackley, 1997) but is not explicitly discussed in WG1 AR4 is simplicity – which is
1224 directly tied to the number and complexity of model assumptions. Yet WG1 AR4
1225 does not, recall, tell us how to map combinations of model virtues onto non-
1226 comparative assessments of model confidence. It tells us when confidence should be
1227 increased on the basis of model virtues but not when confidence should be high. The
1228 IBE approach does and does so in a way that aims to take structural inadequacy into
1229 account.

1230 Theories and models explain phenomena in the sense that they provide
1231 derivations or simulations that show how phenomena are caused or fit into broader
1232 patterns of phenomena (Bokulich, 2011). Thus, GCMs can be said to explain GMST
1233 trends and rising sea levels because the simulations they provide show how these
1234 phenomena causally depend on anthropogenic greenhouse gas trends. How good the
1235 explanations of a model or theory are depends on what combination of virtues the
1236 model or theory has. How good a climate model's explanations are, for example,
1237 depends on how accurate its simulations are, how detailed its descriptions of climatic
1238 mechanisms are, the extent to which it can simulate climate in different periods and so
1239 on. This allows proponents of the IBE approach to propose that how confident we
1240 should be in a theory or model depends on how good the explanations it provides are,
1241 and thus on how good its virtues make its explanations (Lipton, 2004; Thagard, 1978).

1242 That is, it allows the proposal that IBE determines how confident we should be in our
1243 explanations. IBE, as applied to models, is just that form of inference which involves:

- 1244 (i) the possession of alternative explanations of a body of data, where
1245 each alternative explanation rests on a model that explains the data;
- 1246 (ii) a determination of which of the available alternative models that
1247 explain the data provides the best available explanation of the data, i.e.,
1248 of which of these models has the best combination of explanatory
1249 virtues;
- 1250 (iii) an inference to the approximate truth of that model which provides the
1251 best available explanation, provided that the model explains the data
1252 well enough (this standard presentation of IBE has been adapted from
1253 Katzav (2012)).

1254 Since very successful theories do turn out to suffer from unexpected
1255 imperfections, even the most optimistic proponents of the IBE approach only allow
1256 that the very best explanations are good enough. Explanations that are good enough
1257 are usually identified with explanations that are not only empirically successful,
1258 simple, of wide scope and well grounded in background knowledge but that also
1259 provide confirmed novel predictions, that is confirmed predictions of phenomena that
1260 were out-of-sample when they were made *and* unexpected at the time. The idea
1261 behind this stringent definition is that, while it is true that the history of science
1262 provides examples of successful theories and models that have turned out to be
1263 fundamentally wrong, those theories or models which generate confirmed novel
1264 predictions arguably tend to survive, at least as approximations, in later theories (see
1265 Psillos (1999, pp. 101-111) for a standard discussion). Newtonian mechanics is one of
1266 the most successful theories ever, and it lead to its share of novel and confirmed

1267 predictions. Of course, like the already mentioned Newtonian Earth-Sun models,,
1268 Newtonian mechanics appears to be fundamentally wrong in many ways. But
1269 Newtonian mechanics can still be argued to be approximately true. After all, general
1270 relativity does show that we can recover the central equations of Newtonian
1271 mechanics given the right approximations.

1272 Unfortunately, IBE does not provide a way of assessing the quality of specific
1273 classes of CMPs from climate model successes. The IBE approach, like the
1274 confidence building approach in WG1 AR4, provides a way of establishing
1275 confidence in models as wholes (Katzav, 2012).

1276 Further, how accurate a climate model is depends not only on how good its
1277 explanations are but also on how well its parameterization schemes have been
1278 engineered to compensate for our limited ability to model climate. So confidence in a
1279 climate model, or in its fitness for some purpose, should not depend solely on the
1280 quality of its explanations (Katzav, 2012). As to the question whether, in any case,
1281 climate models' explanations are good enough to warrant inferring their approximate
1282 correctness, it is too complex to be addressed here.

1283 We also need to note the dispute about whether IBE should be relied on. When
1284 asked why we should think that IBE allows us to infer the approximate correctness of
1285 models when the future might provide us with surprises about model imperfection,
1286 proponents of IBE answer that we can only explain the success of our models by
1287 supposing that they are approximately true. The success of models would, otherwise,
1288 be a miracle (see, e.g., Musgrave (1988) and Worrall (2010)). Winsberg, however,
1289 provides examples of highly successful principles that do not appear to be
1290 approximately true (Winsberg, 2006). Opponents of IBE point out, further, that the
1291 justification of IBE is itself a kind of IBE and thus begs the question of whether IBE

1292 is acceptable (Laudan, 1981). The justification aims to get us to trust IBE on the
1293 grounds that the best explanation for the successes of a model is its approximate truth.
1294 Some, partly in light of the circular justification of IBE, aim to eschew IBE all
1295 together. Others, accepting that IBE cannot future proof our estimates of how good
1296 our models are, weaken IBE so that it is a form of inference that allows us to rank
1297 models according to explanatory capacity but that leaves open the question of how our
1298 best models relate to the truth. Yet others insist that IBE is fine roughly as it is,
1299 arguing that it is impossible, on pain of an infinite regress, to provide non-circular
1300 justification of all basic inferential principles and that IBE is a good candidate
1301 fundamental principle for justifying models and theories (see Psillos (1999) for a
1302 discussion of some of these views).

1303

1304 **8.4 Severe testing, climate models and climate model projections**

1305 The remaining approach to assessing scientific claims that we will discuss is the
1306 severe testing approach. The idea behind the severe testing approach is that the
1307 deliberate search for error is the way to get to the truth. Thus, on this approach, we
1308 should assess scientific claims on the basis of how well they have withstood severe
1309 testing or probing of their weaknesses (Mayo, 1996; Popper, 2005; Rowbottom,
1310 2011). There are a variety of definitions of ‘severe test’. One prominent definition is
1311 Mayo's (Mayo, 1996; Parker, 2008). It, however, requires that for a test of a claim to
1312 be severe it must be very unlikely that the claim would pass the test if the claim were
1313 false, a requirement that very few tests of climate model fitness for purpose fulfill and
1314 thus which would render the severe testing approach largely unhelpful here. We,
1315 accordingly, explore the usefulness of the main alternative definition, which is
1316 Popper's.

1317 According to Popper, an empirical test of a theory or model is severe to the
1318 extent that background knowledge tells us that it is improbable that the theory or
1319 model will pass the test. Background knowledge consists in established theories or
1320 models other than those being tested (Popper, 2002, p. 150). Popper offers the 1919
1321 test of general relativity's prediction of the precise bending of light in the Sun's
1322 gravitational field as an example of a severe test. The observed bending was
1323 improbable and indeed inexplicable in light of background knowledge at the time,
1324 which basically consisted in Newtonian mechanics. For similar reasons, the precise
1325 precession of Mercury also provided a severe test of general relativity.

1326 A crucial difference between the severe testing approach and the approaches
1327 pursued by WG1 AR4 is that the severe testing approach never allows mere
1328 agreement, or increased agreement, with observations to count in favor of a claim.
1329 That simulation of observed phenomena has been successful does not tell us how
1330 unexpected the data are and thus how severely the data have tested our claims. If, for
1331 example, the successful simulation is the result of tuning, then the success is not
1332 improbable, no severe test has been carried out and no increased confidence in model
1333 fitness for purpose is warranted. Notice, however, that the fact that claims are tested
1334 against in-sample data is not itself supposed to be problematic as long as the data does
1335 severely test the claims [Mayo, 1996]. This is illustrated by the prediction of the
1336 precession of Mercury. The prediction was not novel or even out-of-sample. It was
1337 well measured by Le Verrier in 1859 and was known by Einstein when he constructed
1338 his theory (Earman and Glymour, 1978). Another crucial difference between the
1339 severe testing approach and those pursued by WG1 AR4 is that the severe testing
1340 approach is not probabilistic. The degree to which a set of claims have withstood
1341 severe tests, what Popper calls their degree of corroboration, is not a probability.

1342 How might one apply a (Popperian) severe testing approach to assessing
1343 projection quality? What we need, from a severe testing perspective, is a framework
1344 that assigns a degree of corroboration to a CMP, p , as a function of how well the
1345 model (or ensemble of models), m , which generated p has withstood severe tests of its
1346 fitness for the purpose of doing so. Such severe tests would consist in examining the
1347 performance of some of those of m 's predictions the successes of which would be both
1348 relevant to assessing m 's fitness for the purpose of generating p and improbable in
1349 light of background knowledge. Assessing, for example, a GCM's projection of 21st
1350 century GMST would involve assessing how well the GCM performs at severe tests
1351 of relevant predictions of 20th century climate and/or paleoclimate. That is it would
1352 involve assessing how well the GCM performs at simulating relevant features of the
1353 climate system that we expect will seriously challenge its abilities. A relevant
1354 prediction will be one the accuracy of which is indicative of the accuracy of the
1355 projection of 21st century GMST. Examples of relevant features of the climate the
1356 accurate simulation of which will be a challenge to IPCC-AR5 models are the effects
1357 of strong ENSO events on the GMST, effects of Atlantic sea surface temperature
1358 variations (associated with the MOC) on the GMST and special aspects of the GMST
1359 such as its late 30s and early 40s positive trends. That these data will challenge IPCC-
1360 AR5 models is suggested by the difficulty CMIP3 models have in adequately
1361 simulating them (Katzav, 2011).

1362 The above ideas about applying the severe testing approach will, as a step
1363 towards their operationalization, be elaborated on somewhat and put more formally. p
1364 is corroborated by data just in case the data are probable in light of p but improbable
1365 in light of background knowledge, B . Symbolically, p is corroborated by data just in
1366 case $P(\text{data} | B) < 0.5$ and $C(p | \text{data}, B)$ satisfies

$$1367 \quad C(p | \text{data}, B) \propto P(\text{data} | p, B) - P(\text{data} | B) > 0 \quad (4)$$

1368 Here $P(\text{data} | p, B)$ is the probability of the data in light of p and B , and $P(\text{data} | B)$ is
 1369 the probability of the data in light of B alone. $C(p | \text{data}, B)$ itself results when the
 1370 right hand side of (1) is normalized so as to yield a result that is between 1 and -1,
 1371 where 1 signifies the highest degree of corroboration and -1 signifies the highest
 1372 degree of falsification (Popper, 1983).

1373 Now, we want to assign a degree of corroboration to p as a function of the
 1374 fitness of m for the purpose of generating p . So one could identify $P(\text{data} | p, B)$ with
 1375 the probability that m gives to data which are relevant to testing m 's fitness for the
 1376 purpose of generating p , that is with $P(\text{data} | q, m)$, where q is m 's prediction about the
 1377 relevant data. One could also identify $P(\text{data} | B)$ with the probability given to the
 1378 relevant data by an established rival, $m1$, to m , that is with $P(\text{data} | q1, m1)$, where $q1$
 1379 is $m1$'s prediction for the data. Thus, in the context of assessing m 's suitability for
 1380 generating p , (4) could be interpreted as:

$$1381 \quad C(p | \text{data}, m, m1) \propto P(\text{data} | q, m) - P(\text{data} | q1, m1) > 0 \quad (5)$$

1382 If one's focus is on assessing IPCC-AR5 projections of 21st century GMST, it is
 1383 natural to identify the probability background knowledge gives to data with the
 1384 probability the CMIP3 ensemble gives to them. Accordingly, one could, for example,
 1385 calculate the degree of corroboration of the projection of GMST of a particular AR5
 1386 GCM for the 21st century in light of the model's simulation of data relating to ENSO
 1387 strength by calculating the difference between the probability the model gives to these
 1388 data – $P(\text{data} | q, m)$ in (5) – and the probability the CMIP3 ensemble gives to them –
 1389 $P(\text{data} | q1, m1)$ in (5).

1390 How might the severe testing approach help us with the difficulties involved in
 1391 assessing projection quality? The severe testing approach allows us to bypass any

1392 worries we might have about tuning since it only counts success that does not result
1393 from tuning, success that surely does exist, in favor of CMPs (Katzav, 2011). The
1394 severe testing approach can thus, at least, be used as a check on the results of
1395 approaches that do not take tuning into account. If, for example, the subjective
1396 Bayesian approach assigns a high probability to a projection and the severe testing
1397 approach gives the projection a high degree of corroboration, we can at least have
1398 some assurance that the probabilistic result is not undermined by tuning.

1399 Underdetermination in choice between parameters/available parameterization
1400 schemes might also be addressed by the severe testing approach. Substituting different
1401 parameterization schemes into a model might result in varying degrees of
1402 corroboration, as might perturbing the model's parameter settings. Where such
1403 variations exist, they allow ranking model fitness for purpose as a function of
1404 parameter settings/parameterization schemes. Similarly, degrees of corroboration can
1405 be used to rank fitness for purpose of models with different structures. The resulting
1406 assessment has, like assessment in terms of real possibilities, the advantage that it is
1407 less demanding than probabilistic assessment or assessment that is in terms of truth or
1408 approximate truth. Ranking two CMPs as to their degrees of corroboration, for
1409 example, only requires comparing the two CMPs. It does not require specifying the
1410 full range of alternatives to the CMPs. Nor does it require that we take some stand on
1411 how close the CMPs are to the truth, and thus that we take a stand on the effects of
1412 unknown structural inadequacy on CMP accuracy. Popper's view is that a ranking in
1413 terms of degrees of corroboration only provides us with a ranking of our conjectures
1414 about the truth. The most highly corroborated claim would thus, on this suggestion, be
1415 our best conjecture about the truth. Being our best conjecture about the truth is, in
1416 principle, compatible with being far from the truth.

1417 Consider now some of the limitations of the severe testing approach. To begin
1418 with, while the fact that the severe testing approach is, in some respects, less
1419 demanding than other approaches has its advantages, it also have its disadvantages.
1420 Suppose we rank a claim according to degree of corroboration. What does this imply
1421 for the usability of the claim in research and in decision making? Popper's suggestion
1422 that the most highly corroborated claim is our best conjecture about the truth suggests
1423 a role for corroboration in the context of research. But when is our best conjecture
1424 close enough to the truth to be relevant to practice, e.g., to decision making (Salmon,
1425 1981)? Popper's response is not straightforward (Miller, 2005). However, one can
1426 make use of Popper's idea that claims should be assessed by severe tests without
1427 buying into the rest of his views about science. The beginnings of an alternative
1428 response is as follows: the overall degree of corroboration of a claim depends not just
1429 on how the claim has done at this or that single test, but also on how broadly it has
1430 been tested. A claim's degree of corroboration is thus correlated with the extent to
1431 which the claim is consistent with the overall way things are and, therefore, with the
1432 extent to which the claim is a real possibility. A high enough degree of corroboration
1433 will, accordingly, allow us to conclude that a claim is a real possibility and that it
1434 should be used in decision making.

1435 Another basic worry is that our description of the severe testing approach
1436 presupposes that we are able to determine, prior to using the severe testing approach,
1437 whether data are relevant to assessing fitness for purpose. This includes sometimes
1438 being able to determine, independently of the severe testing approach, that inaccuracy
1439 in simulating a quantity is not substantially relevant to the accuracy of projections of
1440 other quantities. But being able to provide such determinations is something we
1441 required of adequate approaches to assessing projection quality.

1442

1443 **9. Conclusion**

1444 There remain substantial difficulties for WG1 AR4's (climate-model-based)
1445 approaches to assessing projection quality, particularly because they aim at
1446 probabilistic assessment. Indeed, worries about probabilistic assessment of projection
1447 quality are increasingly being raised by those working on projection quality
1448 assessment (Parker, 2010; Smith, 2006; Stainforth et al., 2007a).

1449 The commonly used versions of the subjective Bayesian approach leave us,
1450 because of their limited ability to represent known climate model imperfection, with a
1451 puzzle about why Bayesian updating should be used. Rougier's version does allow a
1452 more complete representation of model imperfection, though it does not actually
1453 provide us with a way of assessing such imperfection. The likelihood approach was
1454 only briefly discussed. It is limited to assessment that takes uncertainty about
1455 parameter choice into account. The confidence building approach has the advantage
1456 of flexibility. It can, since confidence need not be expressed probabilistically, provide
1457 non-probabilistic assessments. So too, the argumentation it uses can in principle be
1458 extended to providing assessments of fitness for purpose, though it currently tends to
1459 stop at assessing models as such.

1460 In examining approaches not used in WG1 AR4, we saw that the similarity
1461 between the confidence building and IBE approaches suggests that IBE might be used
1462 to extend the confidence building approach. The many who do not share in the
1463 skepticism about IBE will be tempted to use the criterion of explanatory goodness in
1464 order to establish the approximate correctness of climate models. At the same time,
1465 we saw that the IBE approach does not help us to select which CMPs we are entitled
1466 to be confident in. We also saw that considering explanatory quality alone is not the

1467 appropriate way of assessing climate model performance. The critical approach
1468 provides not so much a way of assessing projection quality as one of systematizing
1469 such assessments and legitimizing its results. The legitimization it would provide,
1470 however, is problematic because of the lack of agreement about how to assess
1471 projection quality and because of the need to address the question of when consensus
1472 is a guide to truth.

1473 The possibilist and severe testing approaches are promising in that they
1474 propose specific non-probabilistic measures of CMP quality. The severe testing
1475 approach has the additional advantage that it provides a way of trying to get a handle
1476 on the effects of tuning on CMP accuracy. As we have noted, however, both
1477 possibilist and severe testing approaches face problems.

1478 Some of the difficulties that arise in assessing projection quality are
1479 difficulties that would arise irrespective of actual projection accuracy. Tuning may
1480 well not affect the ability of models reliably to generate some important class of
1481 projections. But our uncertainty about the very practice of tuning means that, even if
1482 the projections in question are accurate and reliably generated, we will find it difficult
1483 to decide whether they are accurate. Similarly, the non-linear nature of the climate
1484 system may well not adversely affect the accuracy of some class of projections. But
1485 our uncertainty about whether non-linearity is pertinent to the projections will mean
1486 that we will find it difficult to decide whether they are accurate. This is frustrating, but
1487 does not alter the predicament we find ourselves in with respect to developing
1488 adequate approaches to assessing projection quality.

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1883

1884 **Captions**

1885 *Fig. 1 Temperature changes relative to the corresponding average for 1901-1950*
 1886 *(°C) from decade to decade from 1906 to 2005 over the entire globe, global land area*
 1887 *and the global ocean. The black line indicates observed temperature change, while*
 1888 *the colored bands show the combined range covered by 90% of recent model*
 1889 *simulations. Red indicates simulations that include natural and human factors, while*
 1890 *blue indicates simulations that include only natural factors. Dashed black lines*
 1891 *indicate decades and continental regions for which there are substantially fewer*
 1892 *observations. Adapted from Hegerl et al., FAQ9.2, Fig. 1 (2007, p. 703).*

1893

1894 *Fig. 2 Projected 21st century global mean temperatures changes for various*
 1895 *greenhouse gas emission scenarios. Solid lines are multi-model global averages of*
 1896 *surface warming for scenarios A2, A1B and B1, shown as continuations of the 20th-*
 1897 *century simulations. These projections also take into account emissions of short-lived*
 1898 *GHGs and aerosols. The pink line is not a scenario, but is for Atmosphere-Ocean*
 1899 *General Circulation Model (AOGCM) simulations where atmospheric concentrations*
 1900 *are held constant at year 2000 values. The bars at the right of the figure indicate the*
 1901 *best estimate (solid line within each bar) and the likely range assessed for the six*
 1902 *SRES marker scenarios at 2090-2099. All temperatures are relative to the period*
 1903 *1980-1999. Adapted from the Synthesis Report for IPCC AR4, Fig. 3.2 (2007, p. 7).*

1904

1905 *Fig. 3 Evolution of the MOC at 30°N in simulations with the suite of comprehensive*
 1906 *coupled climate models from 1850 to 2100 using 20th Century Climate in Coupled*
 1907 *Models (20C3M) simulations for 1850 to 1999 and the SRES A1B emissions scenario*
 1908 *for 1999 to 2100. Some of the models continue the integration to year 2200 with the*
 1909 *forcing held constant at the values of year 2100. Observationally based estimates of*
 1910 *late-20th century MOC are shown as vertical bars on the left. Adapted from Meehl et*
 1911 *al., Fig. 10.15 (2007b, p. 773), who build on Schmittner et al. (2005).*

1912

1913 *Fig. 4. Root-mean-square (RMS) error of 1980–99 surface temperature (averaged*
 1914 *over space, relative to the 40-year reanalysis of the European Centre of Medium*
 1915 *range Weather Forecast) shown as a function of the number of models included in the*
 1916 *model average. Panel (a) shows the December-January-February period (DJF),*
 1917 *panel (b) the June-July-August (JJA) period. Red dashed lines indicate the range*
 1918 *covered by randomly sampling the models for the subset; the red solid line indicates*
 1919 *the average. The RMS error converges to a constant value that is more than half of*
 1920 *the initial value for one model. The black dashed line is a theoretical RMS. If the*
 1921 *model biases were independent, then the RMS error for a large sample of models*
 1922 *should decrease with the square root of the number of models (dotted). The blue line*
 1923 *results if the models are sorted by how well they agree with DJF and JJA*
 1924 *observations combined, and it indicates that the average of a few good models*
 1925 *outperforms an average of more models with poorer performance. Adapted from*
 1926 *Knutti et al., Figs 3(c) and 3(d) (2010, p. 2744).*

1927

1928 *Fig. 5 Constraints on the radiative forcing from the observed atmospheric and*
 1929 *oceanic warming. Probability density functions (PDF) for the total (anthropogenic*
 1930 *and natural) radiative forcing (a–c) and the indirect aerosol forcing (d–f) in the year*

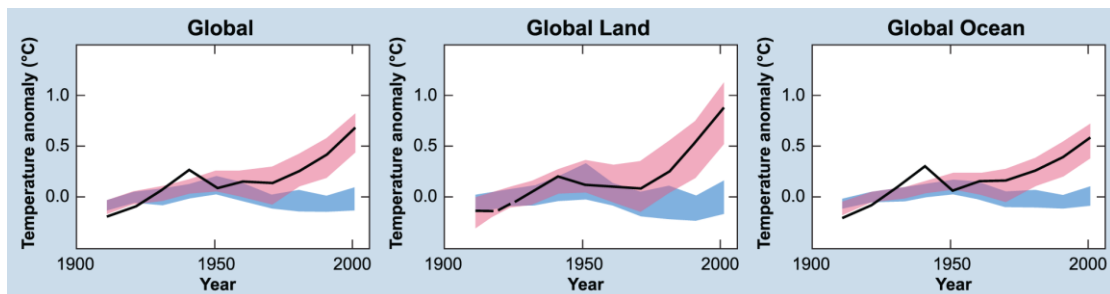
1931 2000 are based on 25,000 Monte Carlo simulations. The initially assumed PDFs are
 1932 given in a and d. The requirement that the model matches the temperature
 1933 observations strongly narrows the PDFs (b and e). If in addition the climate
 1934 sensitivity is restricted to the range adopted by the IPCC (1.5–4.5 K), the PDFs in c
 1935 and f are obtained. Adapted from Knutti et al., Fig. 2 (2002, p. 720).

1936

1937 Fig. 6. The response to parameter perturbations: the frequency distribution of
 1938 simulated climate sensitivity using all model versions (black), all model versions
 1939 except those with perturbations to the cloud-to-rain conversion threshold (red), and
 1940 all model versions except those with perturbations to the entrainment coefficient
 1941 (blue). Adapted from Stainforth et al, Fig. 2(a) (2005, p. 404).

1942

1943 **Figures**

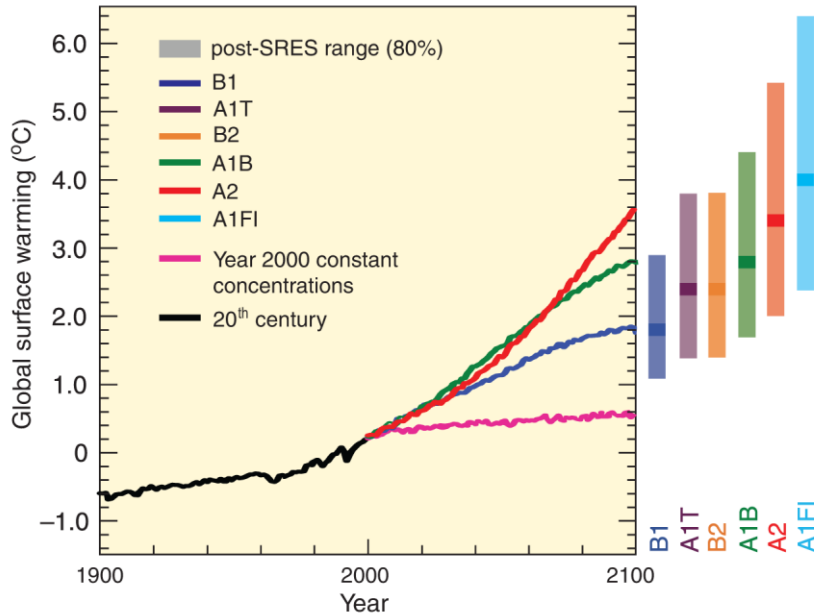


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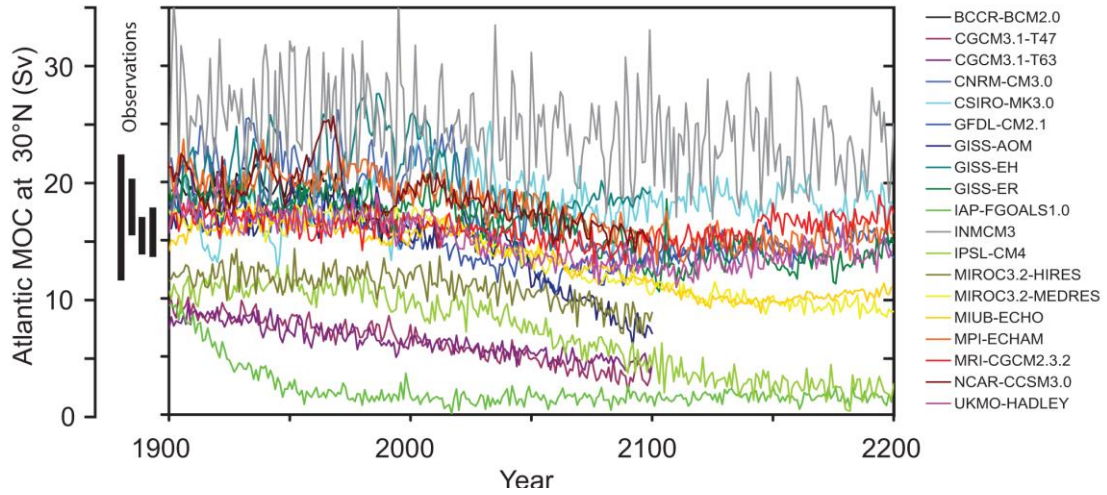
Fig. 1



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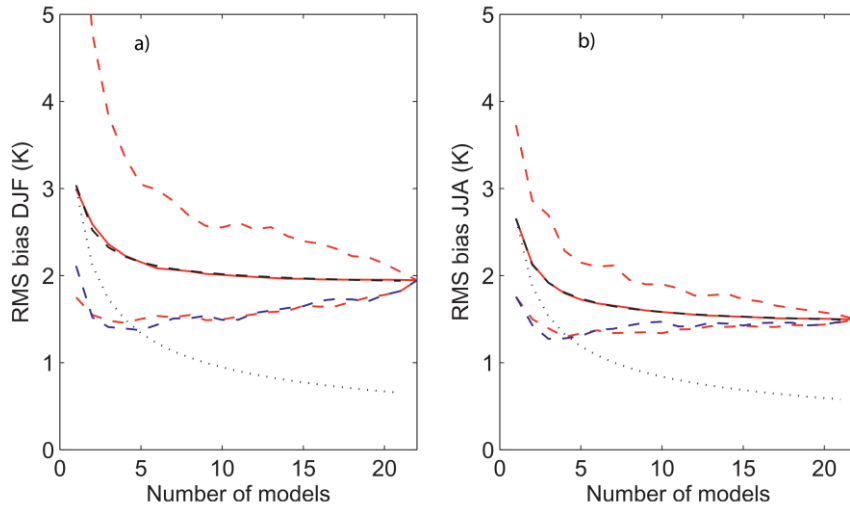
1948 Fig. 2

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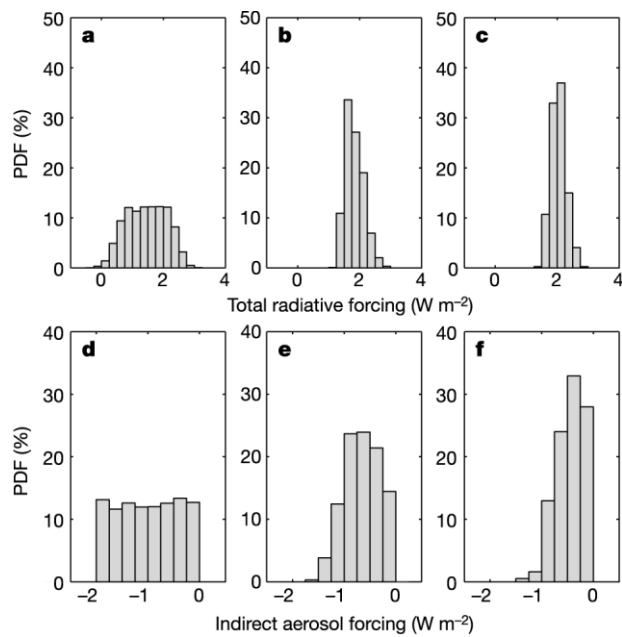
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Fig. 3



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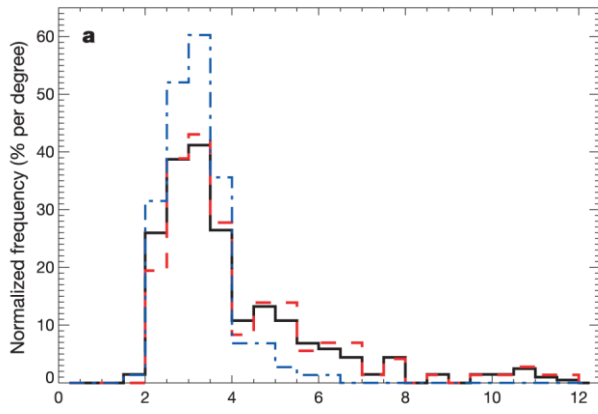
Fig. 4



1956
1957

Fig. 5

1958



1959
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Fig. 6